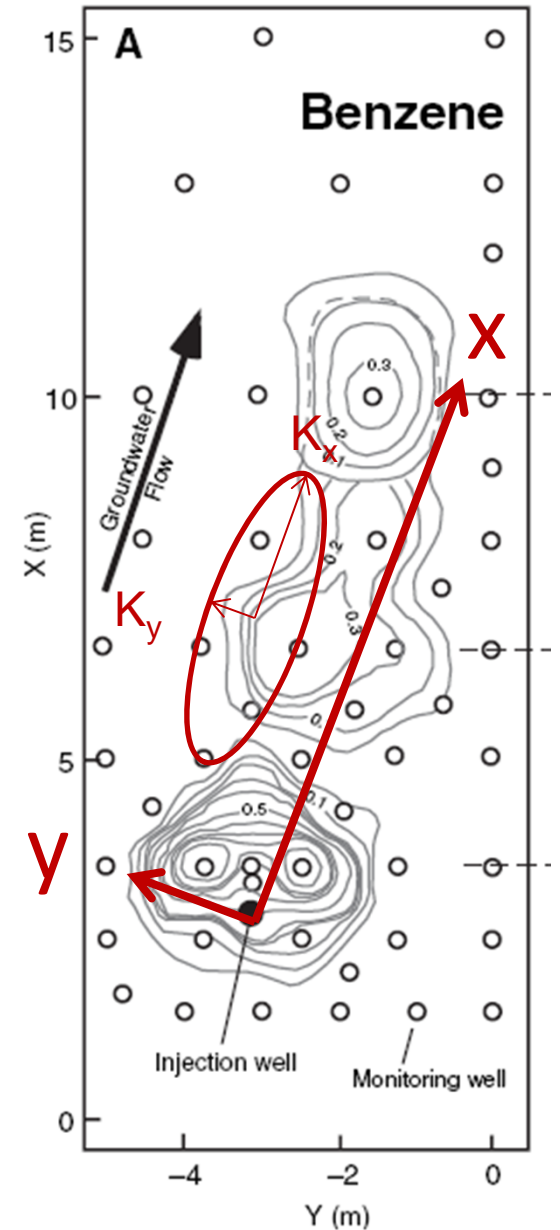


# Gradient Tutorial

# Hydraulic Gradient

*Important to understand for hydrogeological analysis because we use flexible coordinate to study flow and transport in an aquifer:*

- **Natural coordinate:** x: E-W; y: N-S; z: up
- **Analysis coordinate:** x: N30E (x is set parallel to maximum K) ; y: N60W; z: up (elevation) or down (depth)



# Hydraulic Gradient (1D)

Ordinary differential  $\frac{dh}{dx} \approx \frac{\Delta h}{\Delta x} = \frac{h(x + \Delta x) - h(x)}{\Delta x}$

$$\frac{dh}{ds} = [h(s+\Delta s) - h(s)] / \Delta s$$

$$= [h_B - h_A] / \Delta s < 0$$

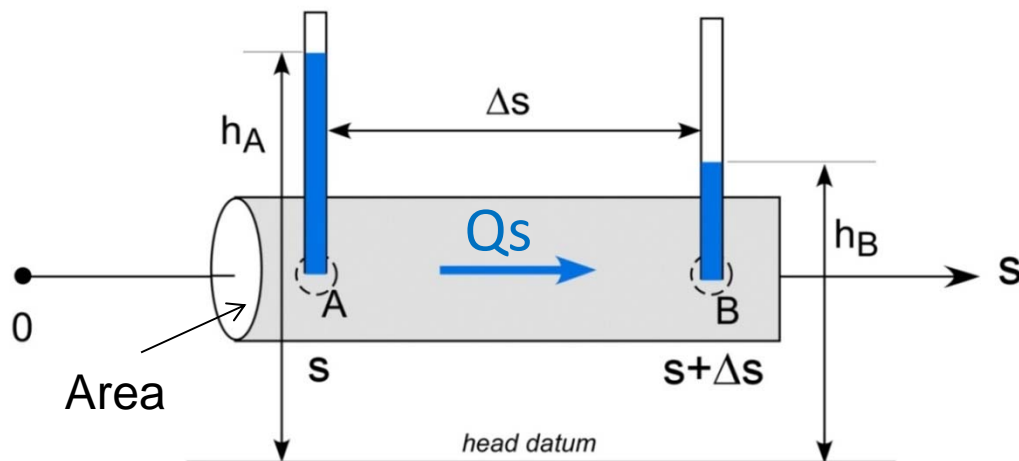
**Darcy's Law**

$$Q_s = -K * \text{Area} * (dh/ds) > 0$$

(correct!  $Q_s$  occurs towards positive  $s$ )

1D Analysis

$$\nabla h = dh/ds = ?$$

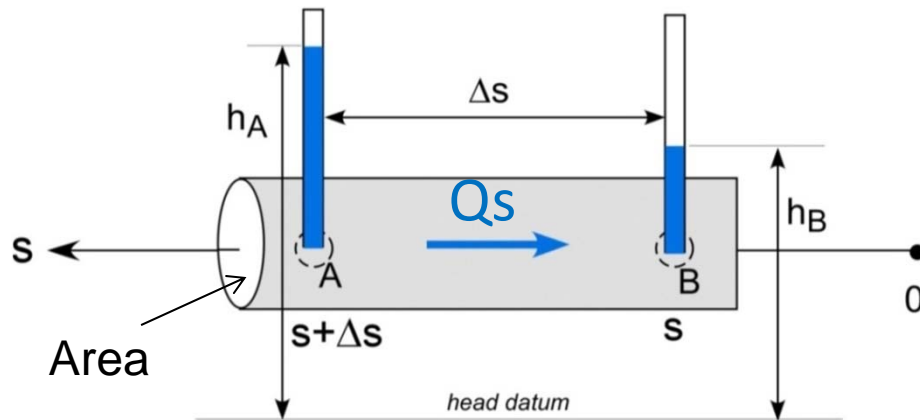


$\nabla$  is the Gradient Operator in 1D:  $\nabla = d[ ]/ds$

$\Delta s$  is a positive distance increment along  $s$  between the two points

## 1D Analysis

$$\nabla h = dh/ds = ?$$



$\nabla$  is the Gradient Operator in 1D:  $\nabla = d[ ]/ds$

$\Delta s$  is a positive distance increment along  $s$  between the two points

When  $s$  points to the opposite direction, the **sign** of gradient ( $dh/ds$ ) and flow rate ( $Q$ ) changes.

$$\begin{aligned} dh/ds &= [h(s+\Delta s) - h(s)]/\Delta s \\ &= [h_A - h_B]/\Delta s > 0 \end{aligned}$$

**Darcy's Law**

$$Q_s = -K * \text{Area} * (dh/ds) < 0 \text{ (correct! since } Q_s \text{ occurs towards } -s)$$

As long as we use the definitions, we'll always get the correct sign of the flow rate *in relation to the coordinate axis*.

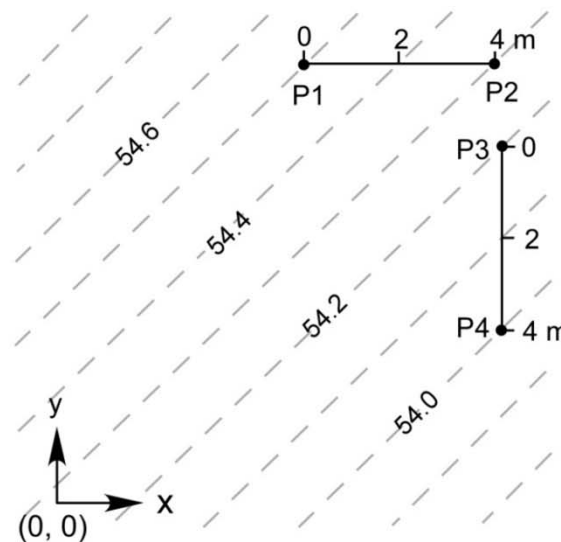
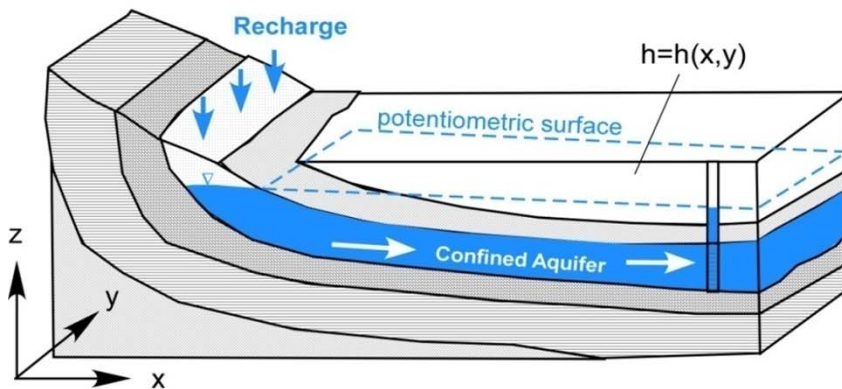
# Hydraulic Gradient (2D)

$$\nabla h = \left\{ \begin{array}{c} \frac{\partial h}{\partial x} \\ \frac{\partial h}{\partial y} \end{array} \right\}$$

$$\frac{\partial h}{\partial x} \Big|_{y \text{ fixed}} \approx \frac{h(x + \Delta x, y) - h(x, y)}{\Delta x}$$

$$\frac{\partial h}{\partial y} \Big|_{x \text{ fixed}} \approx \frac{h(x, y + \Delta y) - h(x, y)}{\Delta y}$$

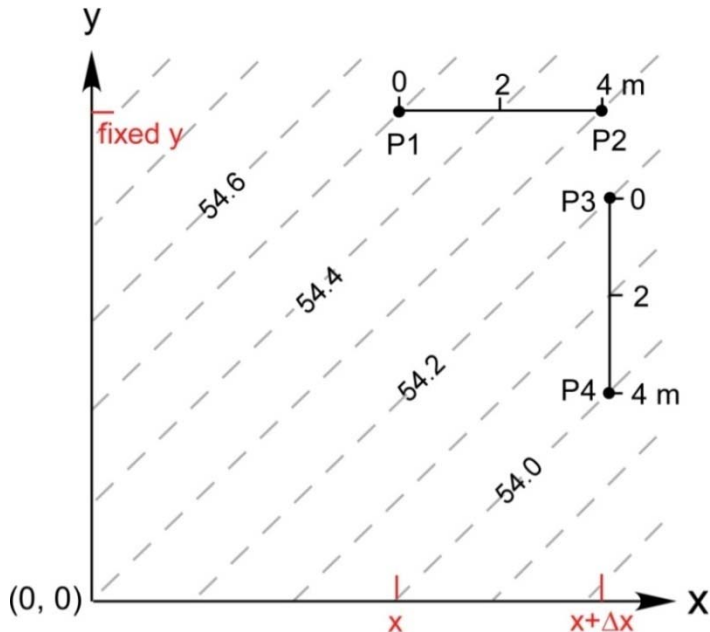
Partial differential



## Mapview

- 4 points where heads are measured: P1, P2, P3, P4;
- Read the heads from the contours

# Hydraulic Gradient (2D)



$$\frac{\partial h}{\partial x} \Big|_{y \text{ fixed}} \approx \frac{h(x + \Delta x, y) - h(x, y)}{\Delta x}$$

$$\frac{\partial h}{\partial y} \Big|_{x \text{ fixed}} \approx \frac{h(x, y + \Delta y) - h(x, y)}{\Delta y}$$

$$\partial h / \partial x = [h(x + \Delta x, y) - h(x, y)] / \Delta x = [h_{P2} - h_{P1}] / \Delta x = -0.05$$

$$\partial h / \partial y = [h(x, y + \Delta y) - h(x, y)] / \Delta y = [h_{P3} - h_{P4}] / \Delta y = +0.05$$

The above results are specific to the **chosen** coordinate;

When the coordinate is changed (e.g., y points S), what happens to  $\partial h / \partial x$  and  $\partial h / \partial y$ ? **Answer:**  $\partial h / \partial x = -0.05$  (x axis is not changed);  
 $\partial h / \partial y = -0.05$  (since P4 occurs at higher y than P3)

Clearly, **SIGN** of the gradient components depends on where the coordinate is pointing (this in turns affects the sign of Darcy flux components).

Evaluate the gradient vector by definition.

Do not use conventional wisdom (“rise over run”). These will not make sense the moment we rotate the coordinate. For example, when z axis is involved, it is common for it to be pointing up (z is elevation) or down (z is depth).

# Hydraulic Gradient (3D)

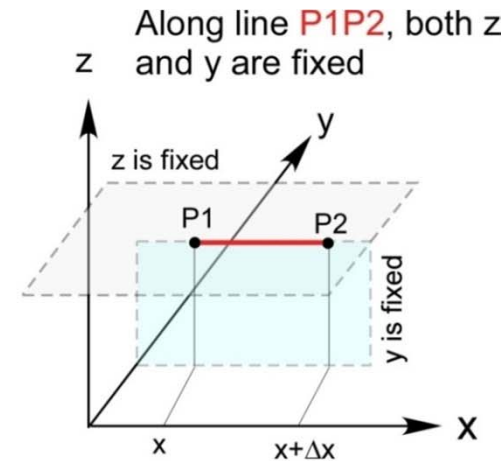
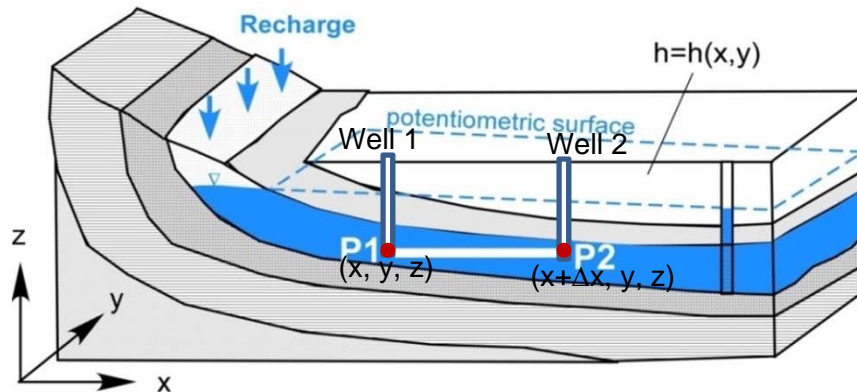
Partial differential

$$\nabla h = \left\{ \begin{array}{c} \frac{\partial h}{\partial x} \\ \frac{\partial h}{\partial y} \\ \frac{\partial h}{\partial z} \end{array} \right\}$$

$$\frac{\partial h}{\partial x} = \lim_{\Delta x \rightarrow 0} \frac{h(x + \Delta x, y, z) - h(x, y, z)}{\Delta x}$$

How the 3D head change with x, fixing y and z

## 3D Gradient Analysis



$P1$  is where well 1 is screened, we call this position  $(x,y,z)$ , we have a head measurement right here at  $(x, y, z)$  (what it is? --water level inside well 1);  $P2$  is where well 2 is screened, we call this position  $(x+\Delta x, y, z)$ , we also have a head measurement right at  $(x+\Delta x, y, z)$  (what it is? --water level inside well 2); We compute:

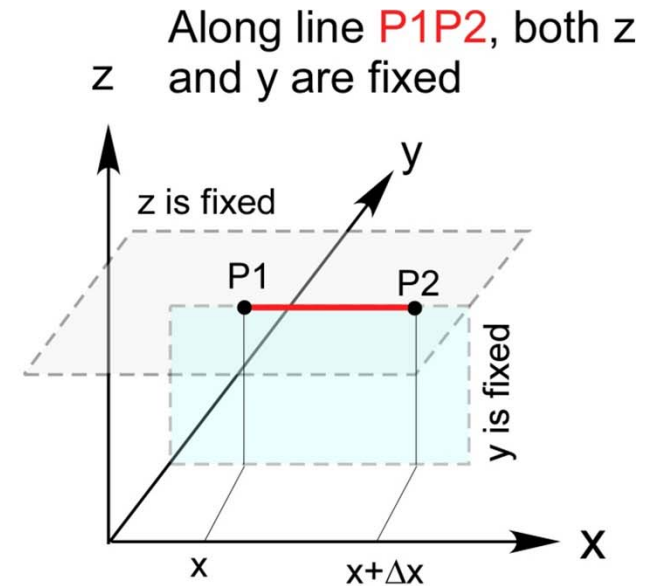
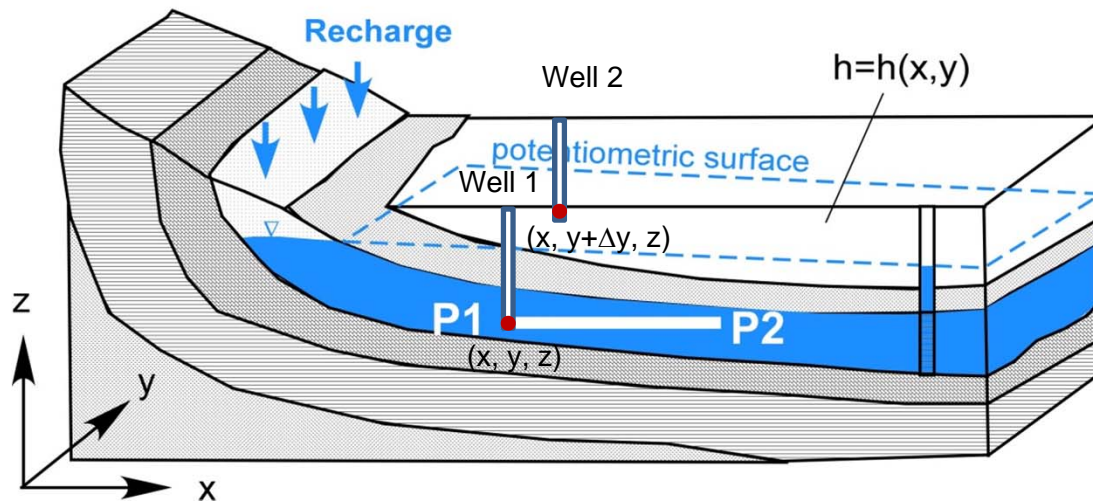
$$\frac{\partial h}{\partial x} = \underbrace{[h(x+\Delta x, y, z) - h(x, y, z)]}_{\text{Head of aquifer measured at aquifer position } (x+\Delta x, y, z)} / \Delta x = [h_{P2} - h_{P1}] / \Delta x = [h_{\text{well2}} - h_{\text{well1}}] / \Delta x \quad (\text{for the given coordinate shown})$$

*P1 and P2 have the same y and z*

Head of aquifer measured at aquifer position  $(x+\Delta x, y, z)$



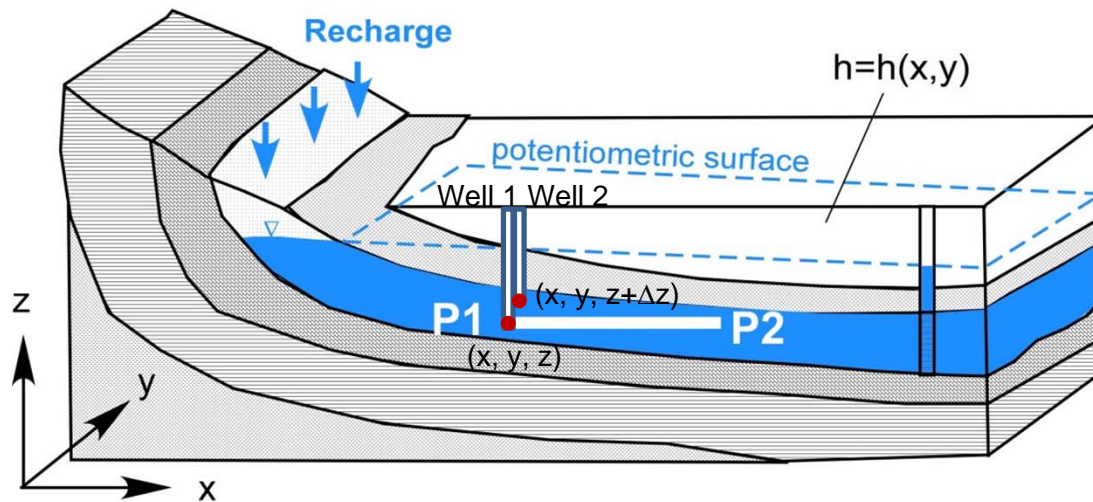
# 3D Gradient Analysis



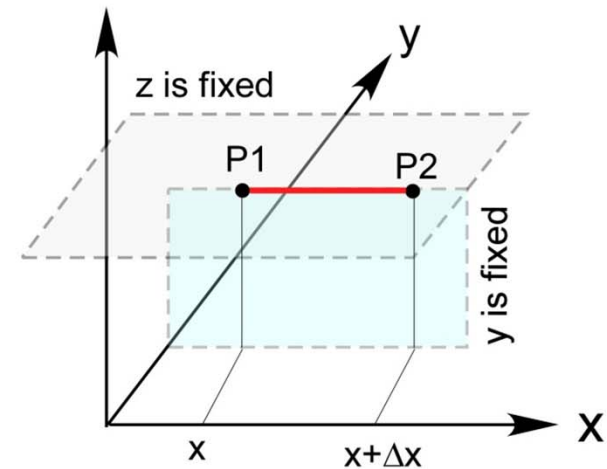
How do we estimate  $\partial h/\partial y$ ?

Well 2 measurement point (screened interval; red dot) will be exactly positioned at  $(x, y+\Delta y, z)$ , while its  $x$ , and  $z$  are the same as the  $x$  and  $z$  of well 1's measurement point (i.e.,  $P1$ ).

# 3D Gradient Analysis



Along line  $P1P2$ , both  $z$  and  $y$  are fixed



## How do we estimate $\partial h/\partial z$ ?

Well 2 measurement point (screened interval; red dot) will be exactly positioned at  $(x, y, z+\Delta z)$ , while its  $x$ , and  $y$  are the same as the  $x$  and  $y$  of well 1's measurement point (i.e.,  $P1$ ).

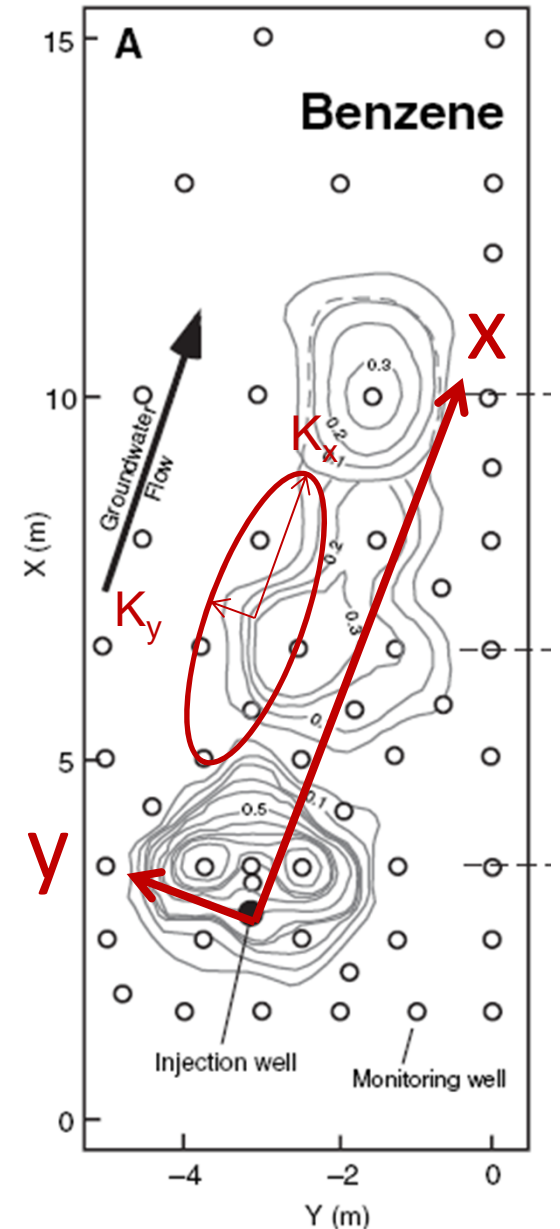
Sometimes instead of drilling a well next to well 1 as shown, we'll use packers to seal off the screened interval at  $P1$ , and perforate the same well casing at a higher elevation to get another reading of the aquifer head at  $(x, y, z+\Delta z)$ .

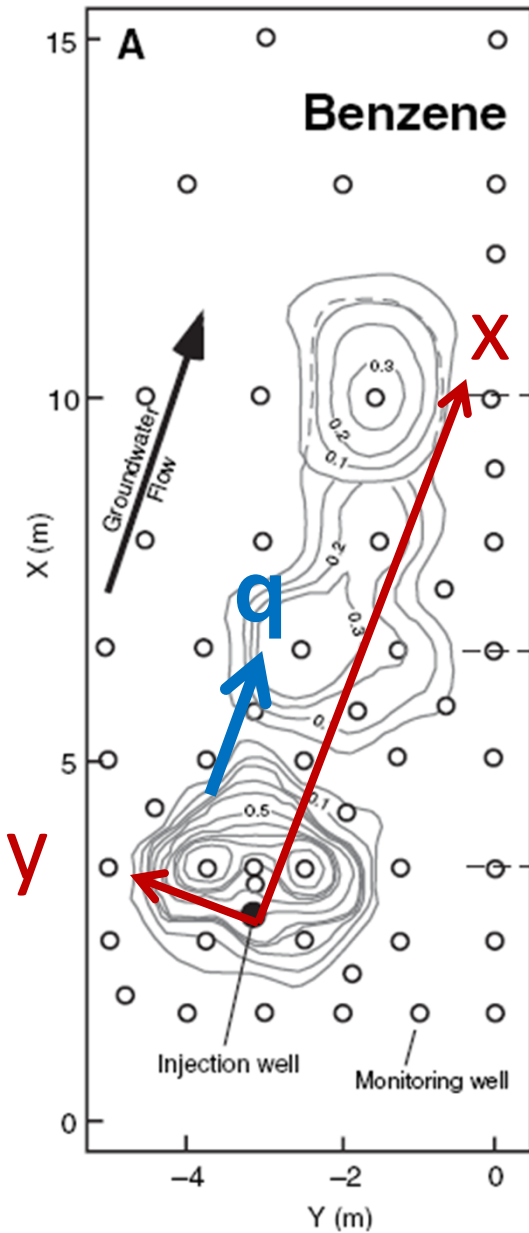
# Let's analyze a 2D mapview problem (vertical flow is ignored: $q_z=0$ )

Here, *analysis coordinate*  $(x,y)$  is chosen because we rotate the axes to be aligned with the principal directions of conductivity, thus  $K^*$  becomes a diagonal tensor:  $K^* = \text{diag}(K_x, K_y, K_z)$ .

Under this condition, we get to use the simplified Darcy's law of this class (*What is it? How do we simplify it to this 2D problem?*)

MODFLOW uses the simplified Darcy's law of this class. The above coordinate rotation is the 1<sup>st</sup> step in MODFLOW.

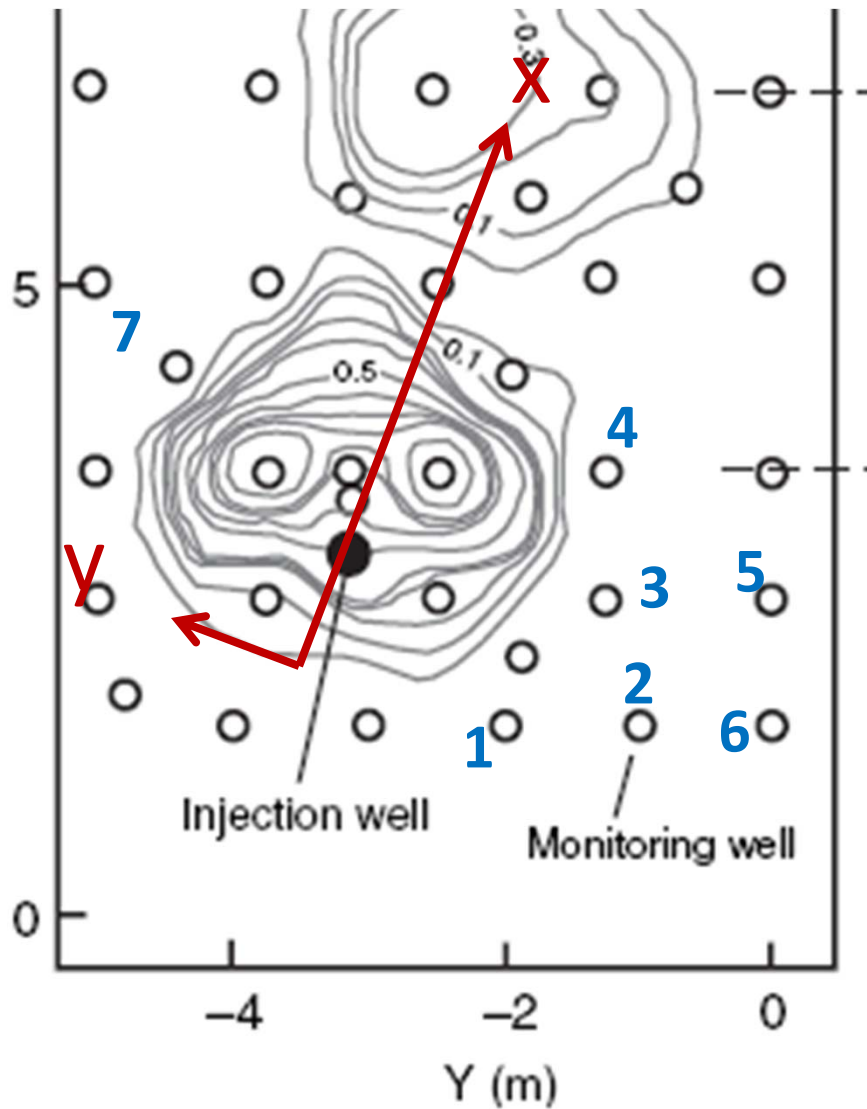




Darcy's flux of the aquifer is known: gw is flowing at **10 m/yr** towards N30E.

- Within the **analysis coordinate** (x,y), x is aligned with N30E, how would you write the Darcy's flux (**q**) vector?

- What will this flux be within the "natural coordinate" (S-E; N-W)?



If  $q$  is not known, we use Darcy's law to find it:

- Within the **analysis coordinate**  $(x,y)$ , how do you use the hydraulic heads from the pizeometers to evaluate  $\partial h/\partial x$ , and  $\partial h/\partial y$ ?

*Hint: where are the wells that satisfy the definitions of  $\partial h/\partial x$ ,  $\partial h/\partial y$ ?*

**Exercise 5** In a 2D transect of the head contours, calculate  $q_x$  and  $q_z$  assuming isotropic conductivity ( $K_x = K_z = 2$  m/day). Make a scaled vector sketch of the  $x$  and  $z$  components of the Darcy flux and the flux vector itself  $\mathbf{q}$  (Assume:  $q_y = 0$ ).

**Repeat** the problem assuming anisotropic conductivity ( $K_x = 2$  m/day,  $K_z = 0.1$  m/day). Discuss how the orientation of  $\mathbf{q}$  relates to the head contours in both cases.

**Here, we assume: principal directions of  $K^*$  are aligned with the coordinate axes; For a 2D transect, the Darcy's law is (please review how 3D Darcy's law is simplified):**

$$\begin{Bmatrix} q_x \\ q_z \end{Bmatrix} = - \begin{bmatrix} K_x & 0 \\ 0 & K_z \end{bmatrix} \begin{Bmatrix} \partial h / \partial x \\ \partial h / \partial z \end{Bmatrix}$$

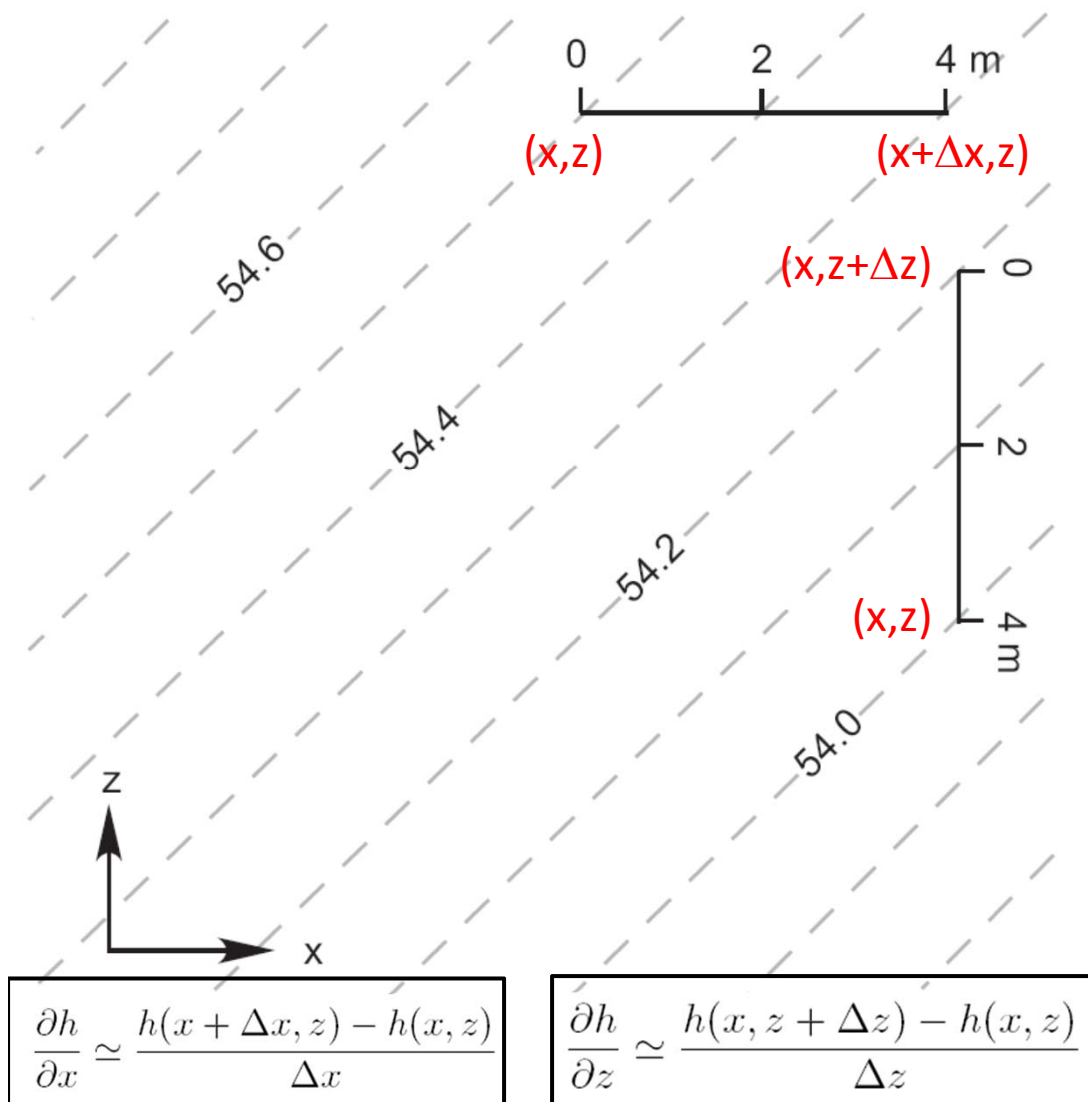
# How would we calculate the directional head gradients in **2D**?

$$\frac{\partial h}{\partial x} \approx \frac{h(x + \Delta x, z) - h(x, z)}{\Delta x} \quad | \text{ fixed } z$$

$$\frac{\partial h}{\partial z} \approx \frac{h(x, z + \Delta z) - h(x, z)}{\Delta z} \quad | \text{ fixed } x$$



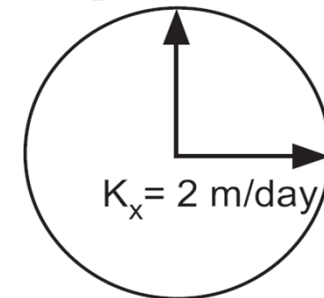
## Head Contours Along a Vertical Transect



$$\begin{Bmatrix} q_x \\ q_z \end{Bmatrix} = - \begin{bmatrix} K_x & 0 \\ 0 & K_z \end{bmatrix} \begin{Bmatrix} \partial h / \partial x \\ \partial h / \partial z \end{Bmatrix}$$

$$\begin{aligned} q_x &= -K_x (\partial h / \partial x) \\ q_z &= -K_z (\partial h / \partial z) \end{aligned}$$

$$K_z = 2 \text{ m/day}$$



Isotropic

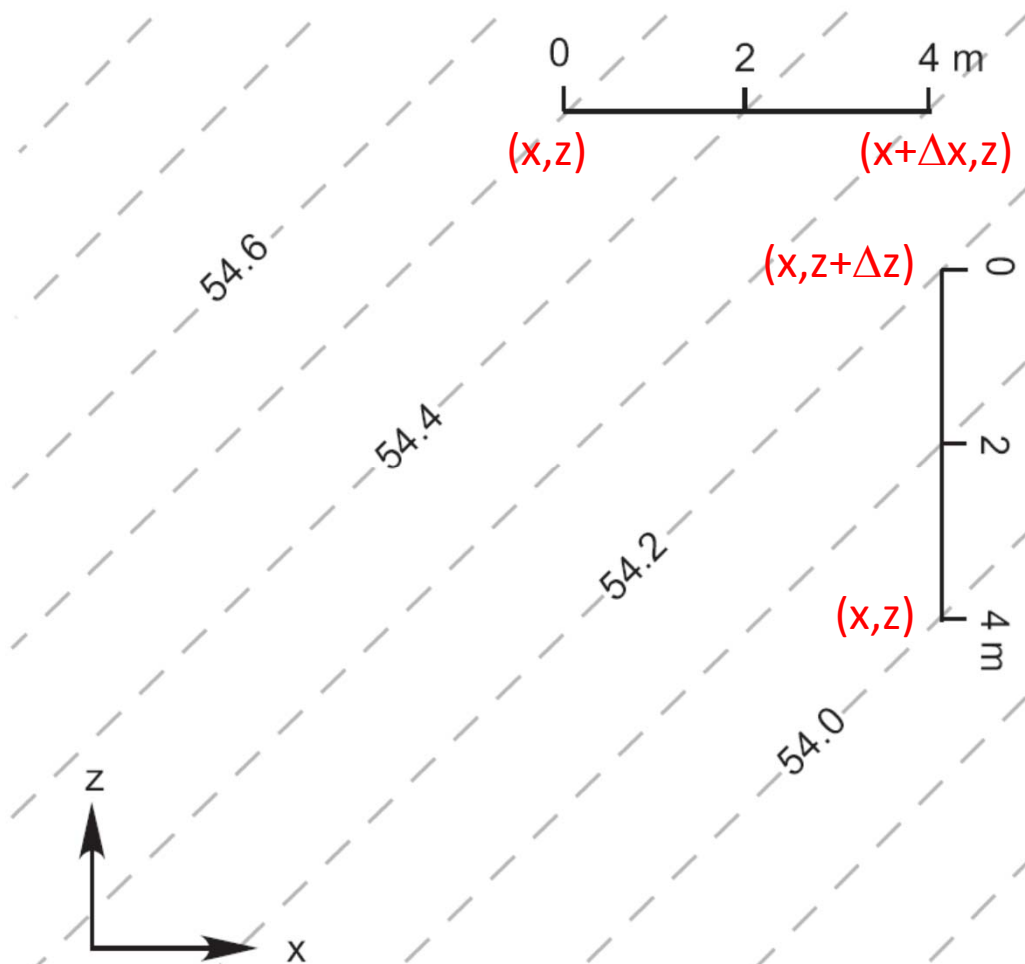
$$\begin{aligned} q_x &= -K (\partial h / \partial x) \\ q_z &= -K (\partial h / \partial z) \end{aligned}$$



## When $K$ is isotropic (a scalar):

- Darcy flux vector is perpendicular to the head contours (or *equipotential lines*):  $\mathbf{q}$  points from higher head towards lower head.
- $\mathbf{q}$  has the **opposite** direction with the head gradient vector:  $\mathbf{l} = \{\partial h / \partial x, \partial h / \partial z\}^T$ ,  $\mathbf{l}$  points from lower head to higher head (*please plot  $\mathbf{l}$  on previous plot*);
- What about the GW streamlines, i.e, macroscopic fluid flow pathways (*please plot on previous plot*)?

## Head Contours Along a Vertical Transect

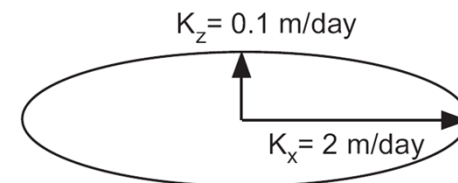


$$\begin{Bmatrix} q_x \\ q_z \end{Bmatrix} = - \begin{bmatrix} K_x & 0 \\ 0 & K_z \end{bmatrix} \begin{Bmatrix} \partial h / \partial x \\ \partial h / \partial z \end{Bmatrix}$$



$$q_x = -K_x \left( \frac{\partial h}{\partial x} \right)$$

$$q_z = -K_z \left( \frac{\partial h}{\partial z} \right)$$

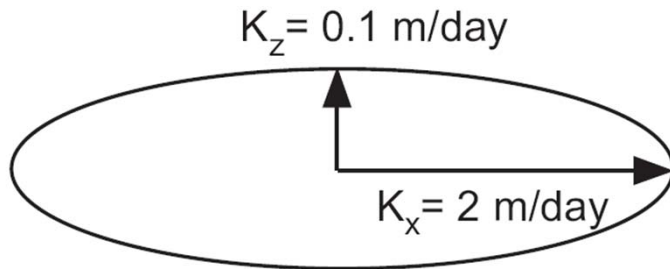


Please plot:  $\mathbf{q}$ ;  $\mathbf{I}$ ; and streamlines

$$\frac{\partial h}{\partial x} \approx \frac{h(x + \Delta x, z) - h(x, z)}{\Delta x}$$

$$\frac{\partial h}{\partial z} \approx \frac{h(x, z + \Delta z) - h(x, z)}{\Delta z}$$

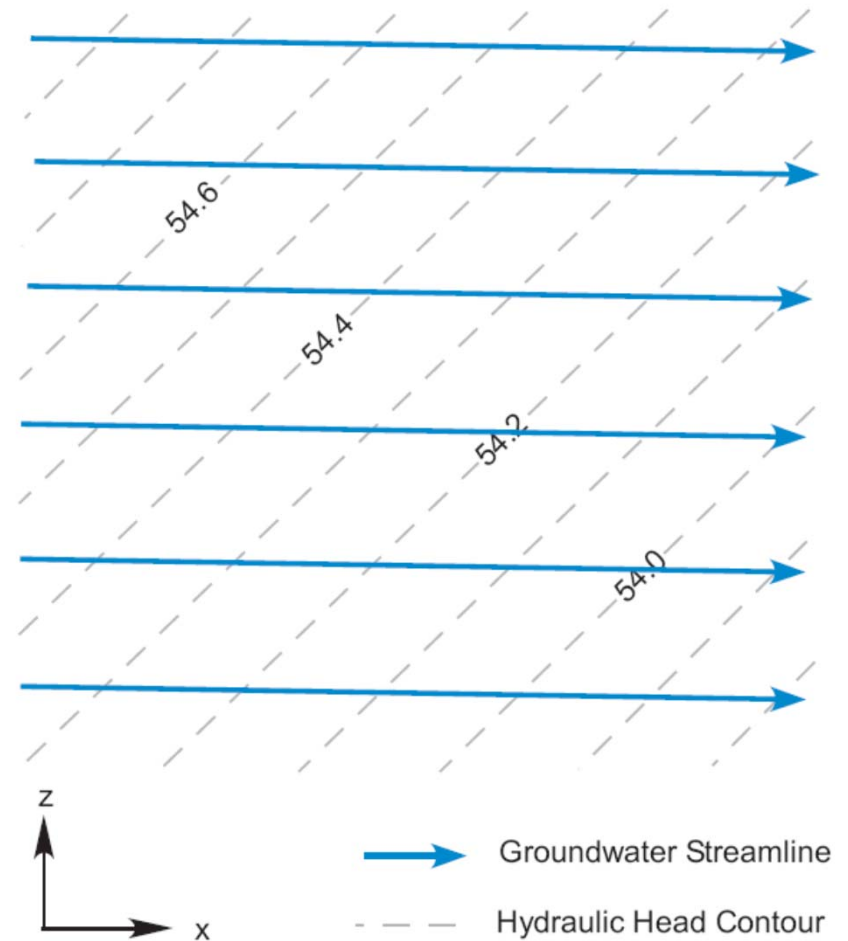
# When K is Anisotropic:



Clearly, though gw always flows from high head to low head, if the K field is not isotropic, gw flow direction will **not** be perpendicular to the head contours---many people make this mistake!

In a field, always think of the possibility of the existence of K anisotropy (**In 4.13 on “Equivalent Conductivity”, we will prove that layered rocks give rise to large scale K anisotropy**)

Vertical Transect (K is Anisotropic)



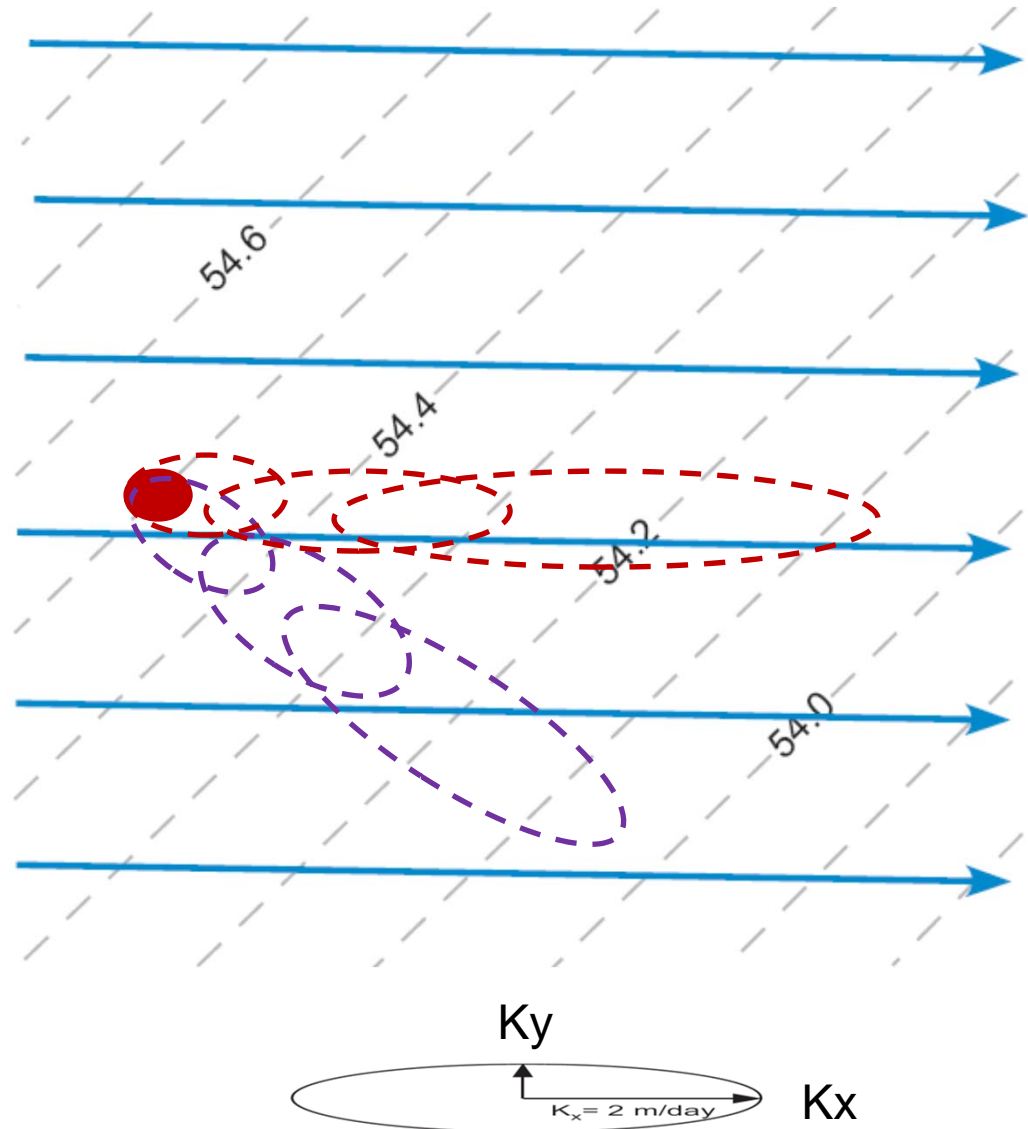
## “Missing Plume”:

GW is contaminated at an undisclosed site;

Engineers first assumed isotropic  $K \rightarrow$  streamline perpendicular to head contours  $\rightarrow$  plume migrates down head gradient; Install wells to intersect the plume for treatment, but no plume is found along the imagined flow path;

Buried “channels” in the aquifer  $\rightarrow K$  is anisotropic  $\rightarrow$  actual flow path is different (blue streamlines)  $\rightarrow$  real plume.

## Plainview



## When K is Anisotropic:

When conductivity is anisotropic ( $K^*$  becomes a tensor), Darcy flux is not perpendicular to the head contours, nor it is in the opposite direction of the head gradient vector.

$$\vec{q} = -K\vec{I};$$

Isotropic ( $K$  is scalar):  $K$  acts to stretch  $\mathbf{I}$  to form  $\mathbf{q}$ ; “-” results in opposite direction of  $\mathbf{q}$  and  $\mathbf{I}$

Anisotropic ( $K^*$  is tensor):  $K^*$  acts to both stretch and rotate  $\mathbf{I}$  to form  $\mathbf{q}$

# In summary, in hydrological analysis:

- (1) Set up your analysis coordinate: *to use the diagonal tensor of  $K^*$  (of this class), coordinate must be aligned with  $K^*$  principle directions;*
- (2) Set head datum somewhere  $\rightarrow$   $z$  & head become defined;
- (3) Within the coordinate, evaluate head gradient components:  $dh/ds$  (1D), or  $\partial h/\partial x$ ,  $\partial h/\partial y$ ,  $\partial h/\partial z$  (2D or 3D; partial sign). *Pick 2 points in the aquifer (usually where the well screens are), find their heads (water level inside wells measuring aquifer heads at the well screens), and calculate the gradient (use definition)*
- (4) Within your coordinate, find Darcy flux components:  $q_x$ ,  $q_y$ ,  $q_z$  which are linked to head gradient by Darcy's Law (1, 2 or 3D form);
- (5) Within your coordinate, we can map out the Darcy flux ( $\mathbf{q}$ ) which describes the *macroscopic* flow field (please review how to plot a vector given its components).