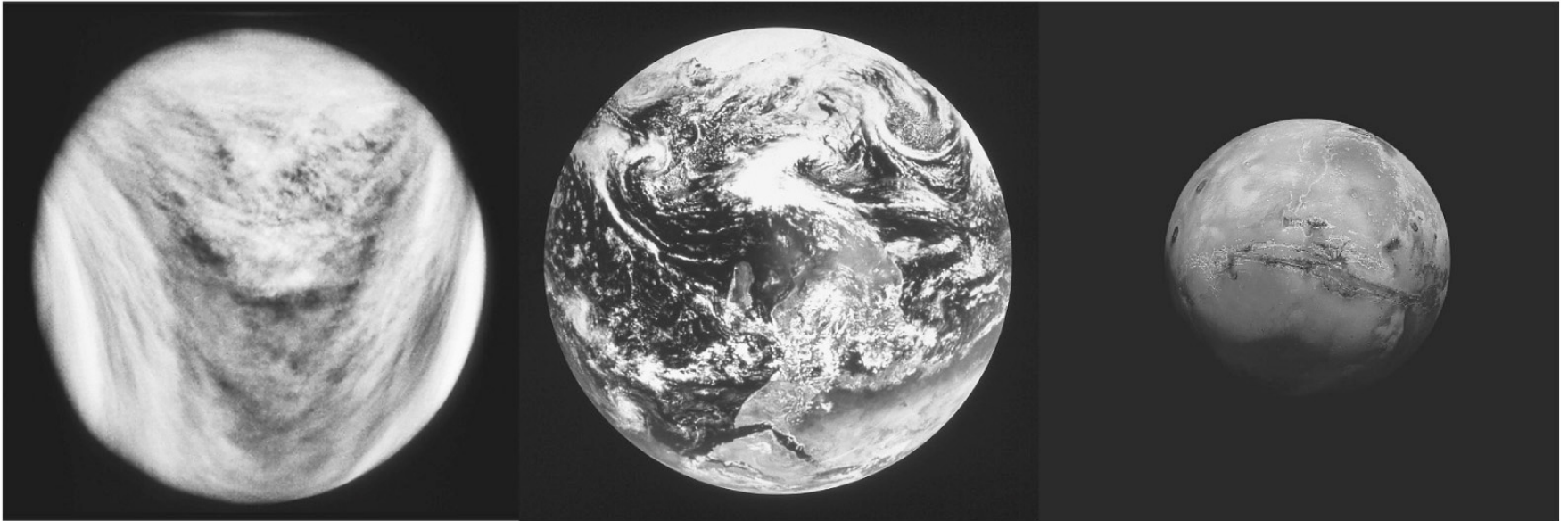


Earth: the Goldilocks Planet



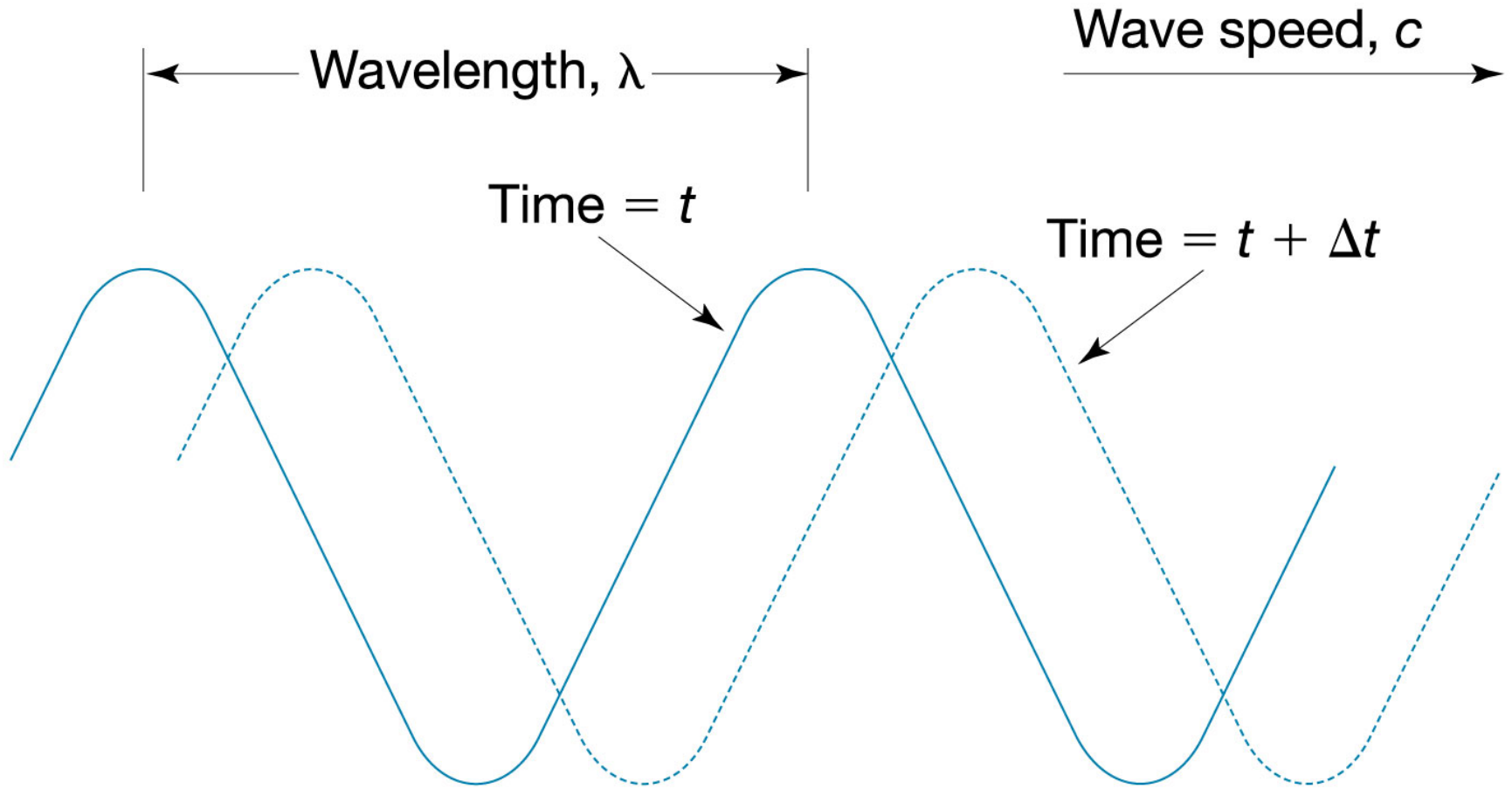
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Not too hot...
(460°C)

Not too cold...
(-55°C)

Fig. 3-1

Wave properties:



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Wavelength, velocity, and ...?

Fig. 3-2

Reviewing units:

Wavelength = distance (meters or nanometers, etc.)

Velocity = distance per unit time (miles per hour, meters per second)

$$\frac{\textit{meters}}{\textit{second}}$$

$$\lambda \nu = c \quad (\text{Meters})(?) = (\text{Meters/second})$$

$$\text{Frequency} = \frac{1}{\textit{time}}$$

$$\text{Example: } \frac{1}{\textit{second}}$$

$$\frac{1}{\textit{second}} = \text{“per second”} = \text{“Hertz”}$$

Energy of *photon*:

$$E = h\nu = hc/\lambda$$

(h = Planck's const.)

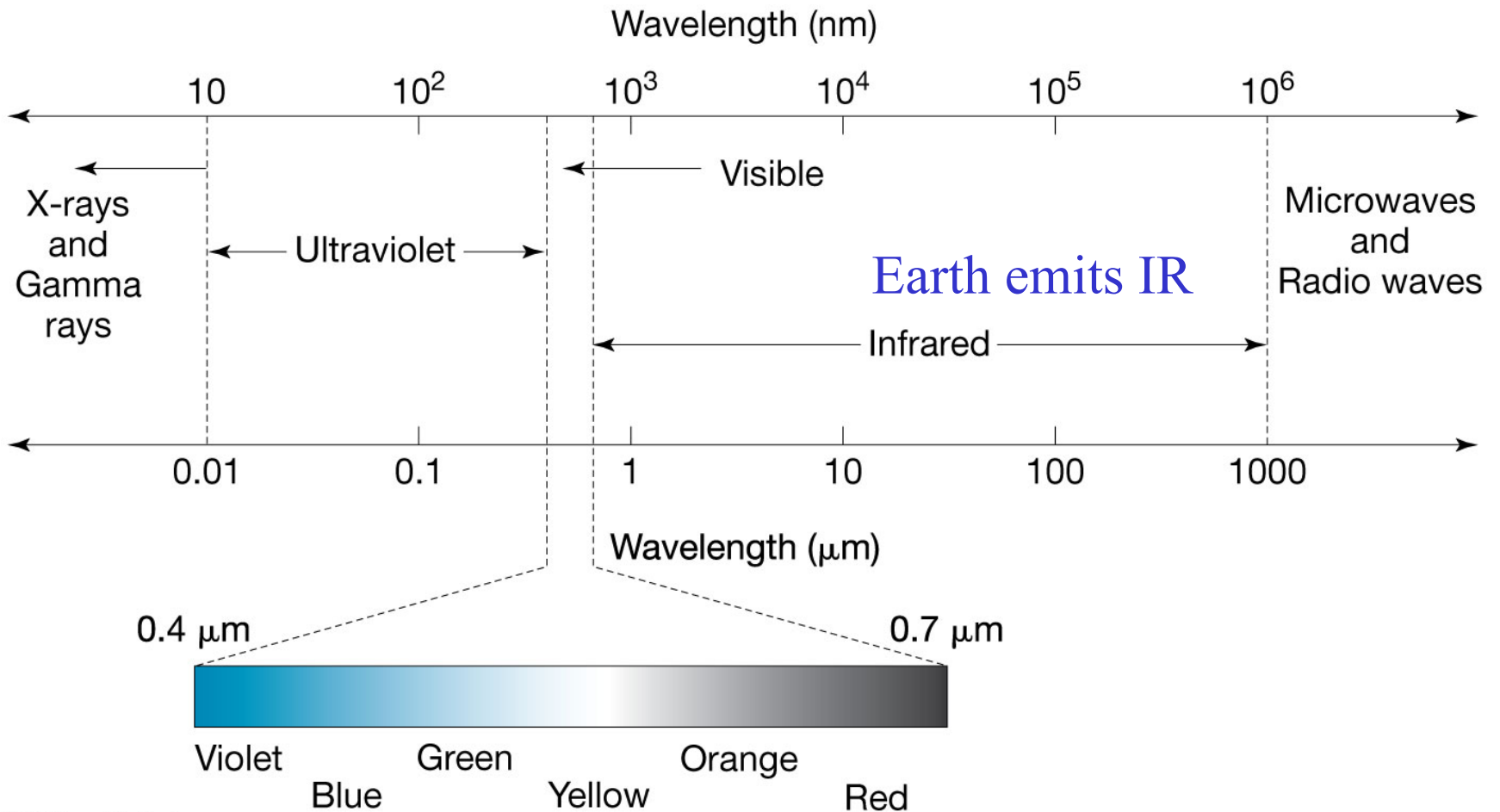
Inverse relationship:

Short λ photons are high E

Long λ photons are low E

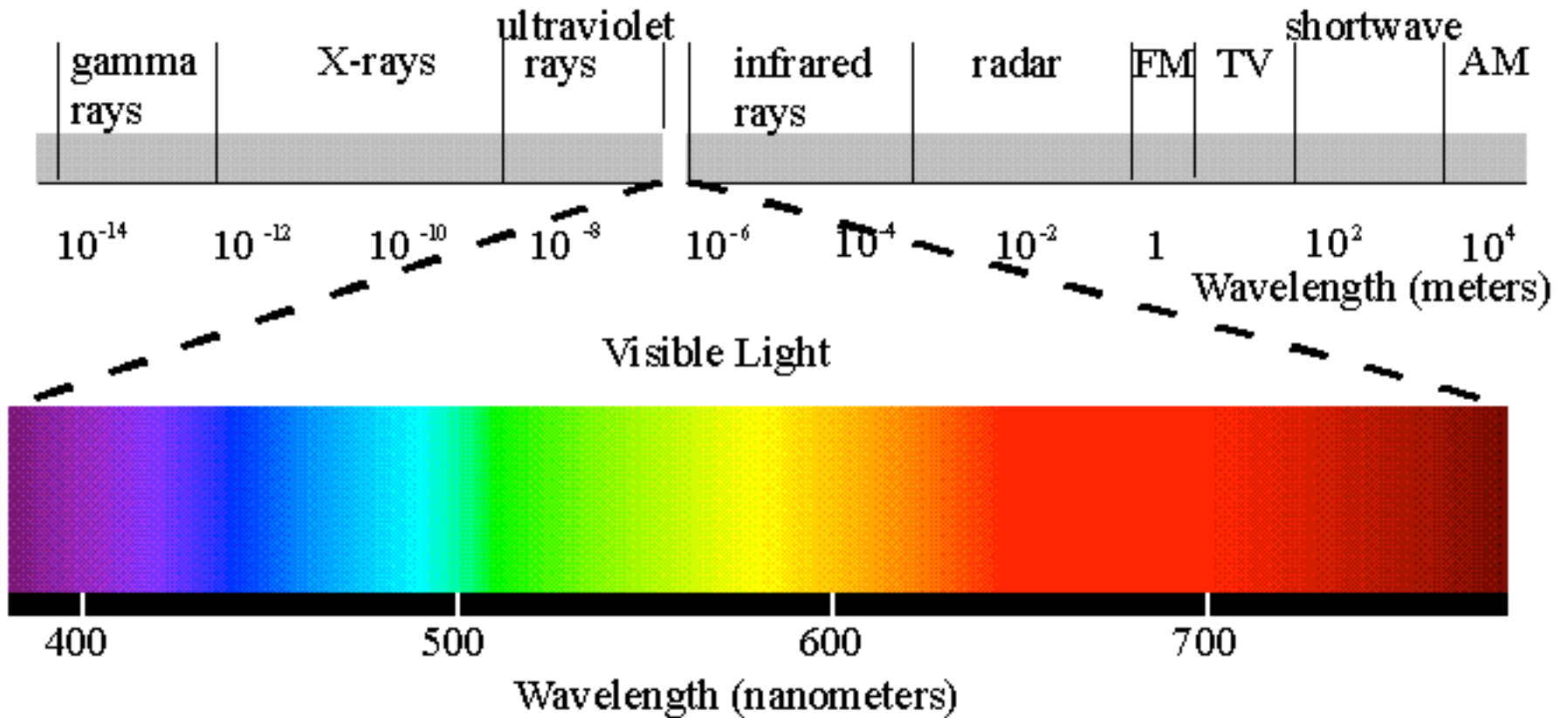
Electromagnetic Radiation:

Fig. 3-3



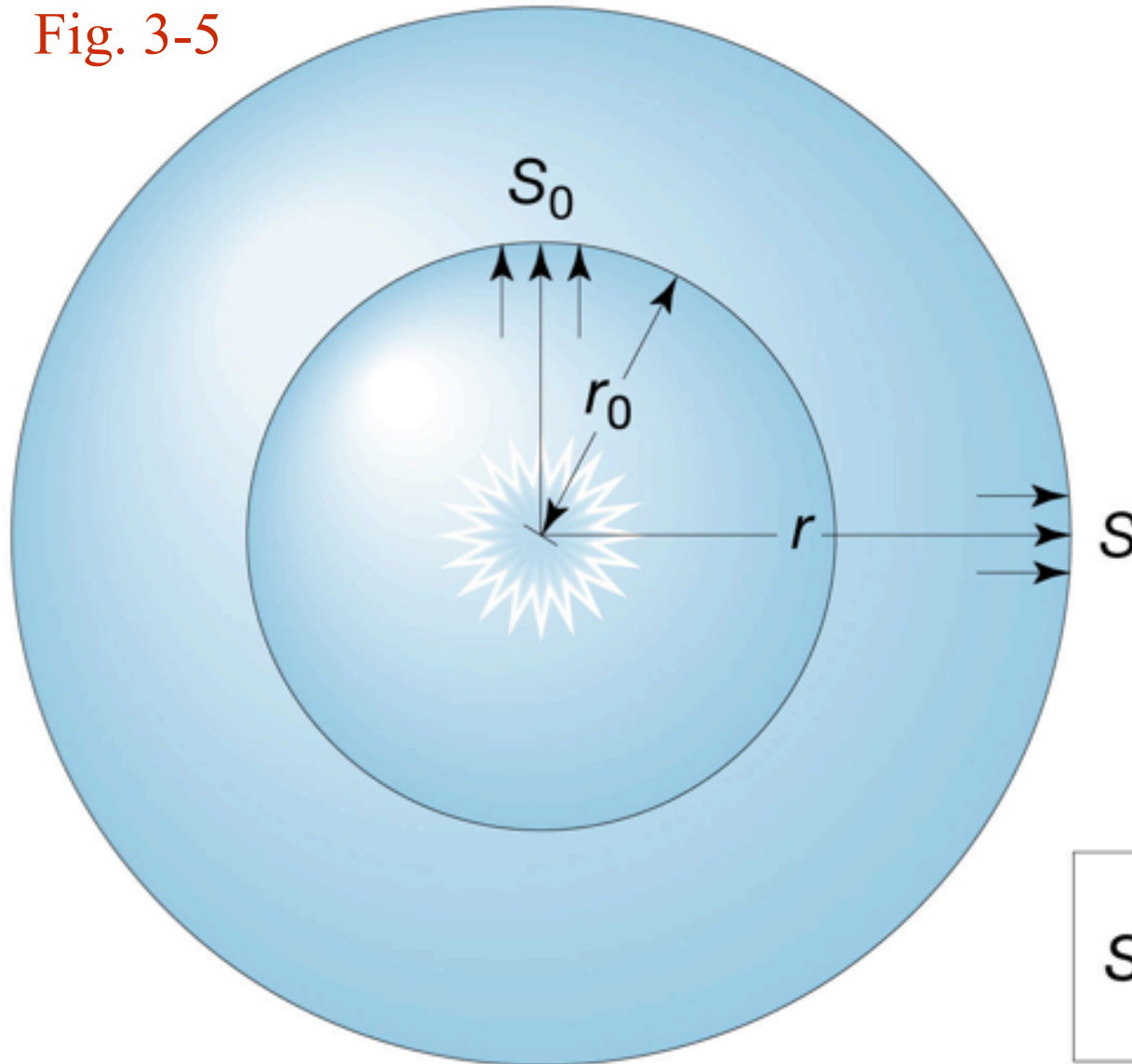
Incoming from Sun

Electromagnetic (“EM”) Radiation:



Flux = energy that passes through a given area per unit time
Units: watts per square meter, **W/m²**

Fig. 3-5



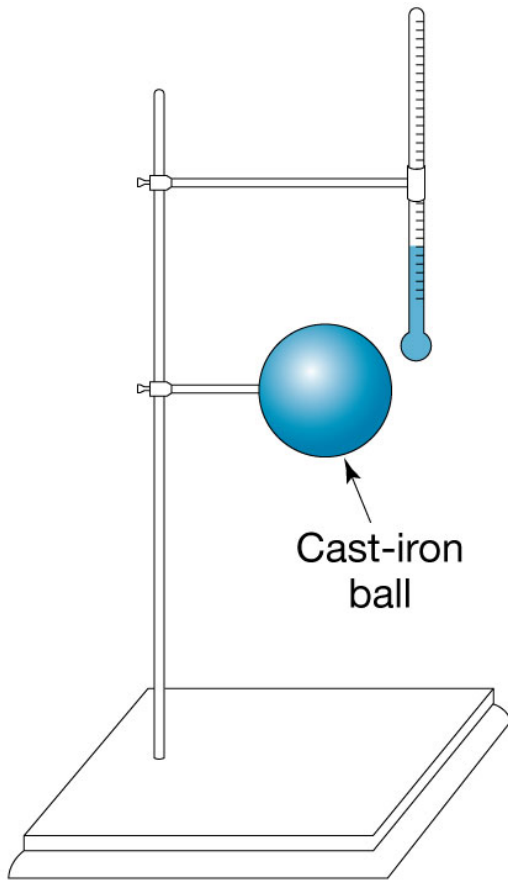
For Sun, energy flux is depends on the distance from the sun, described by inverse-square law:

$$S = S_0 \left(\frac{r_0}{r} \right)^2$$

TABLE 3-1 Freezing and Boiling Points of Water by Temperature Scale

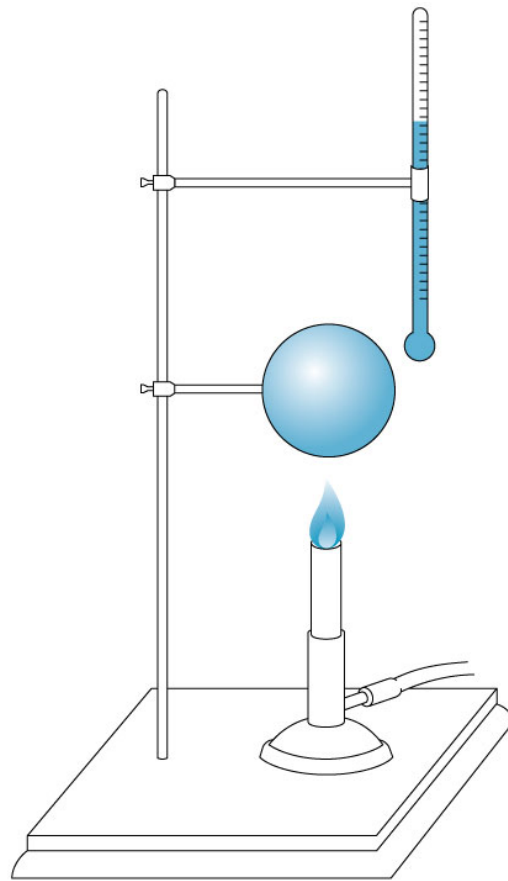
Temperature Scale	Freezing Point	Boiling Point (at sea level)
Fahrenheit	32°	212°
Celsius	0°	100°
Kelvin (absolute)	273.15	373.15

Fig. 3-6

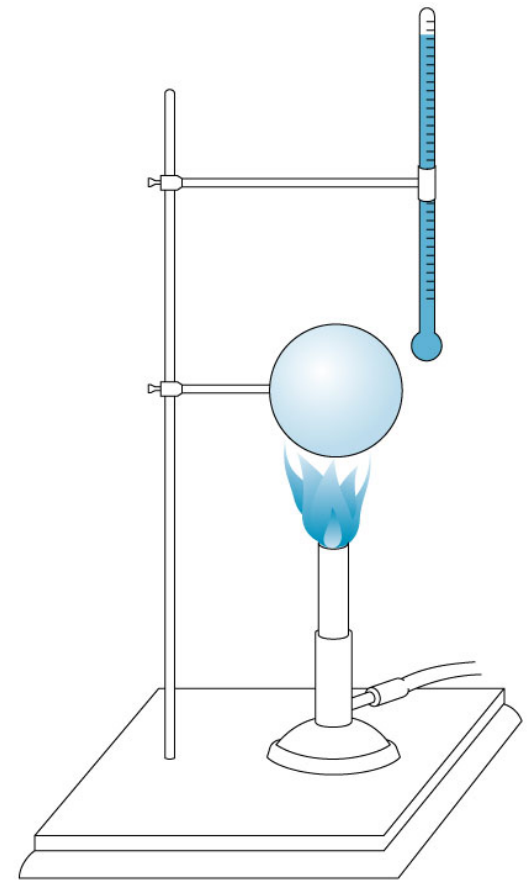


Cast-iron
ball

Room temperature



Hot



Hotter

Temperature varies with color

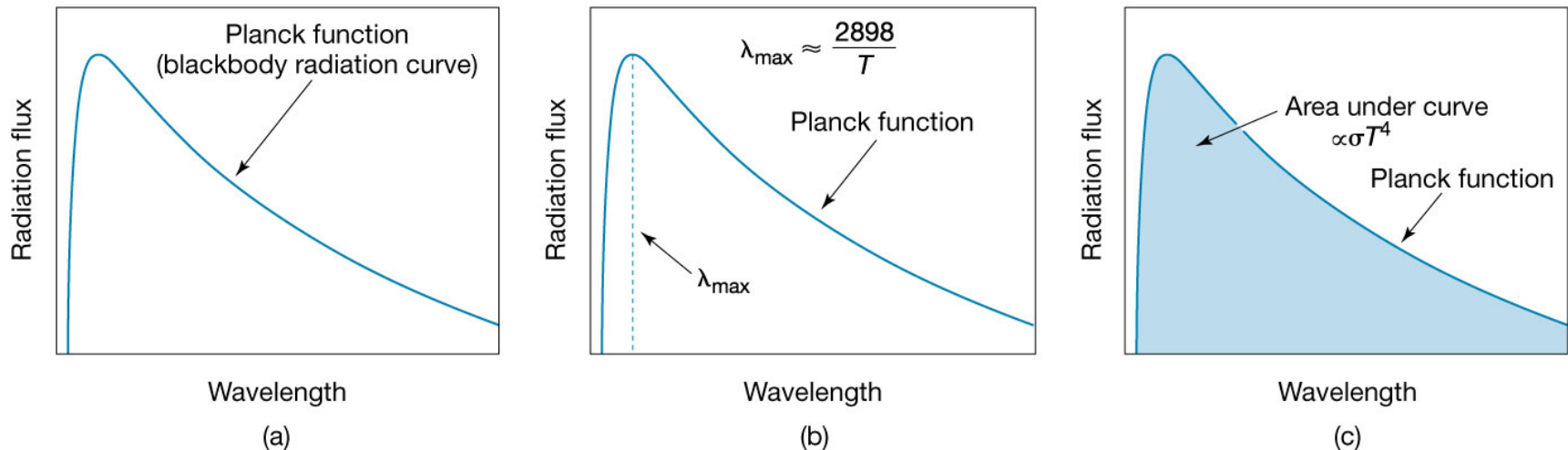


Blackbody radiation:

A blackbody absorbs & emits e-m radiation with 100% efficiency
The wavelength distribution of emitted radiation is a function of T

(red-hot vs. white-hot)

Fig. 3-7 shows a blackbody radiation curve:



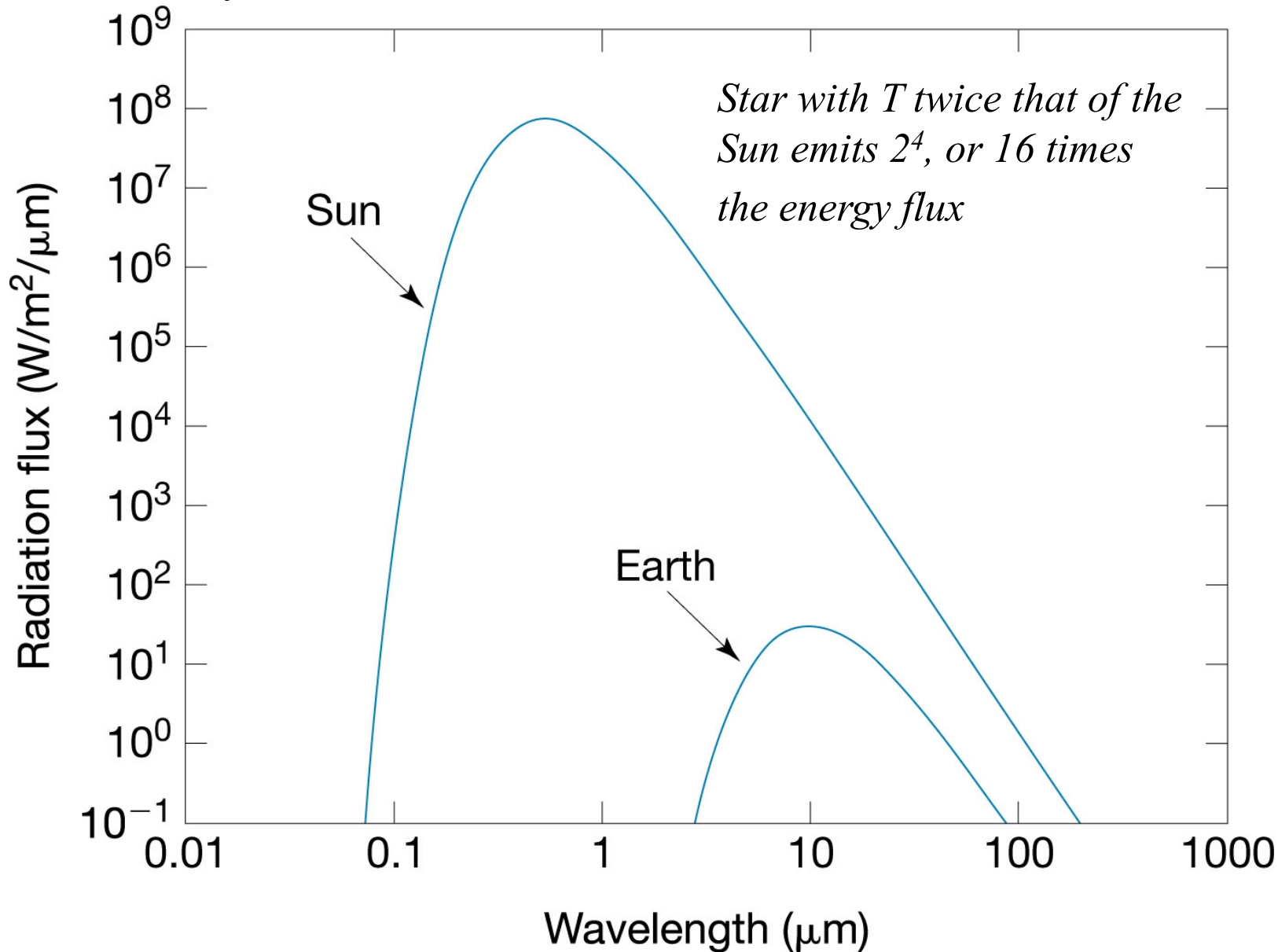
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Wien's law: λ of peak flux is inversely related to body's T

λ_{\max} (in micrometers) = $2898/T$ (in Kelvins)
Sun's T is 5780K; Earth's T is 288K; what is λ_{\max} ?

Stefan-Boltzmann law: total energy flux is proportional to the area under the blackbody radiation curve, and: $F = \sigma T^4$, σ is a constant

Fig. 3-8

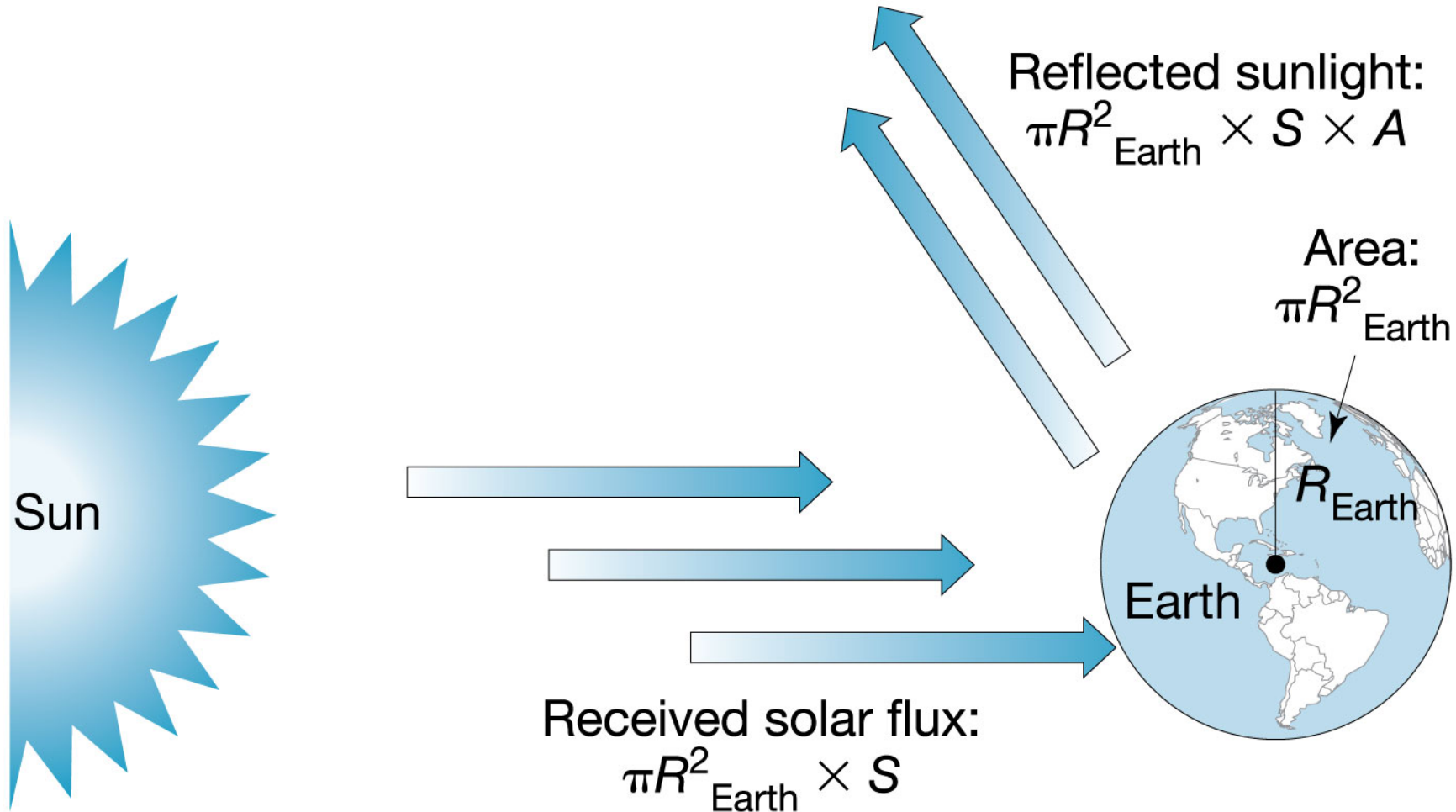


Area of circle: πR^2

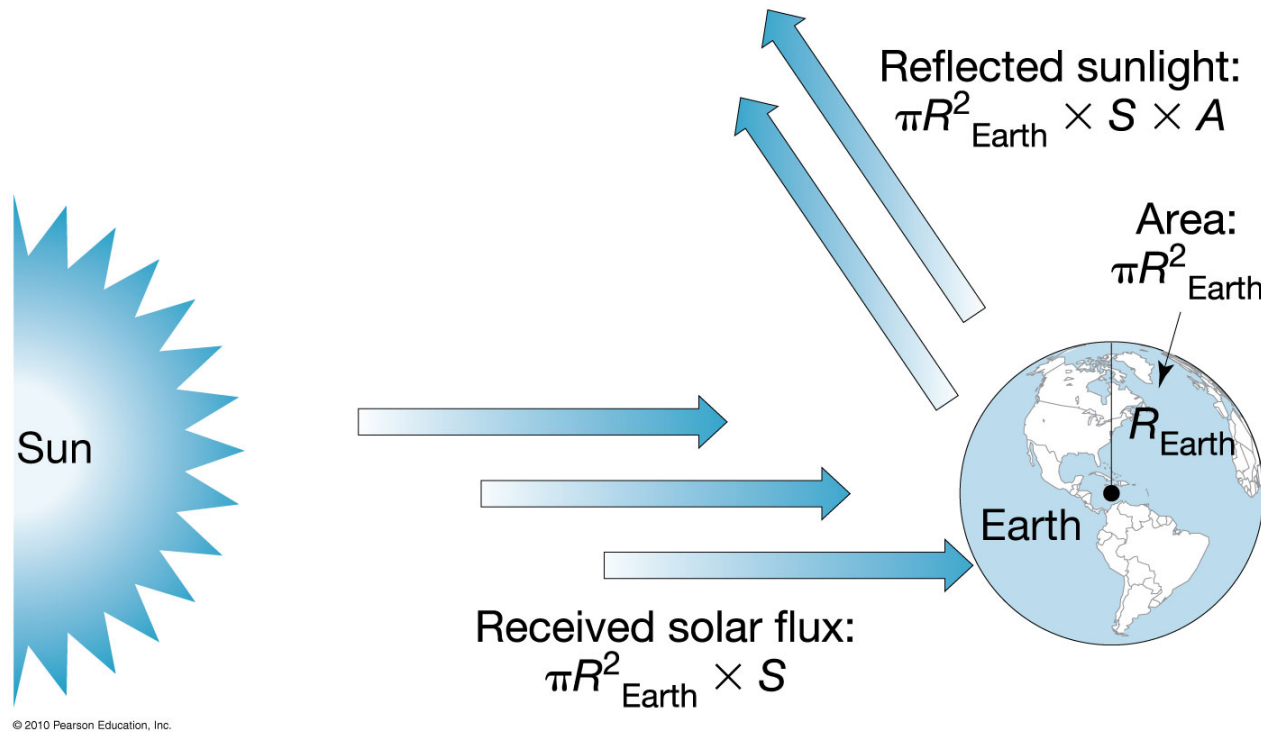
Area of sphere: $4\pi R^2$

Earth's Energy Balance (Budget):

Energy absorbed = Energy emitted



Planetary energy balance: E absorbed = E emitted

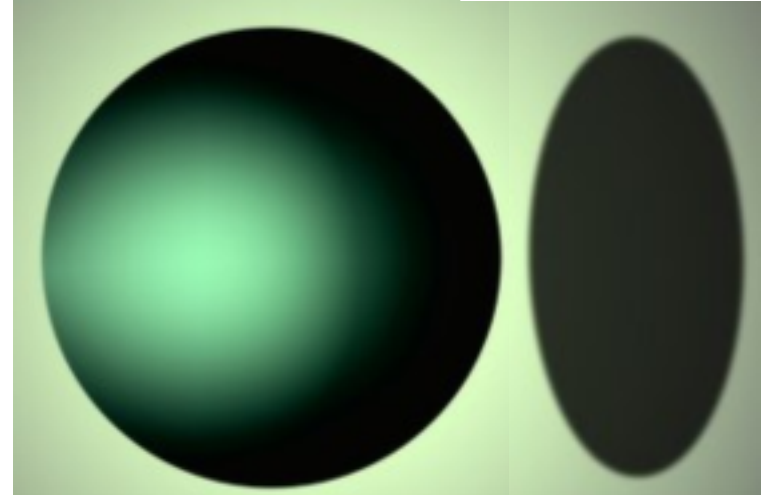
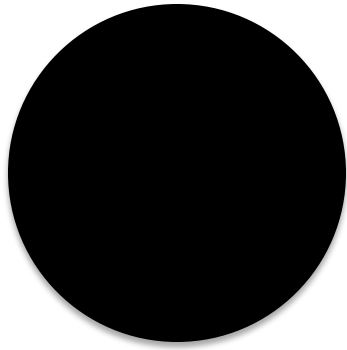


Box Fig. 3-1

S = Power flux at Earth's distance from Sun: $\sim 1,367 \text{ W/m}^2$
What is the power for each square meter of Earth's surface?

Energy absorbed = Energy Intercepted – Energy Reflected

“Intercepted” power ... how many m^2 ?



(E intercepted) - (E reflected) =

$$\mathbf{(\pi R^2 S) - (\pi R^2 S A) = \pi R^2 S (1 - A)}$$

(S = solar flux, A = albedo)

We know energy absorbed (E_{abs}), and from energy balance:

E emitted by Earth = $4\pi R^2 \times \sigma T^4$
(surface area) x (flux per unit area)

$$E_{\text{emit}} = E_{\text{abs}} : 4\pi R^2 \sigma T^4 = \pi R^2 S(1-A)$$

Divide both sides by $4\pi R^2$,

$$\sigma T^4 = (S/4)(1-A)$$

Earth Energy Balance:

$$\sigma T_e^4 = \frac{S}{4} (1 - A)$$

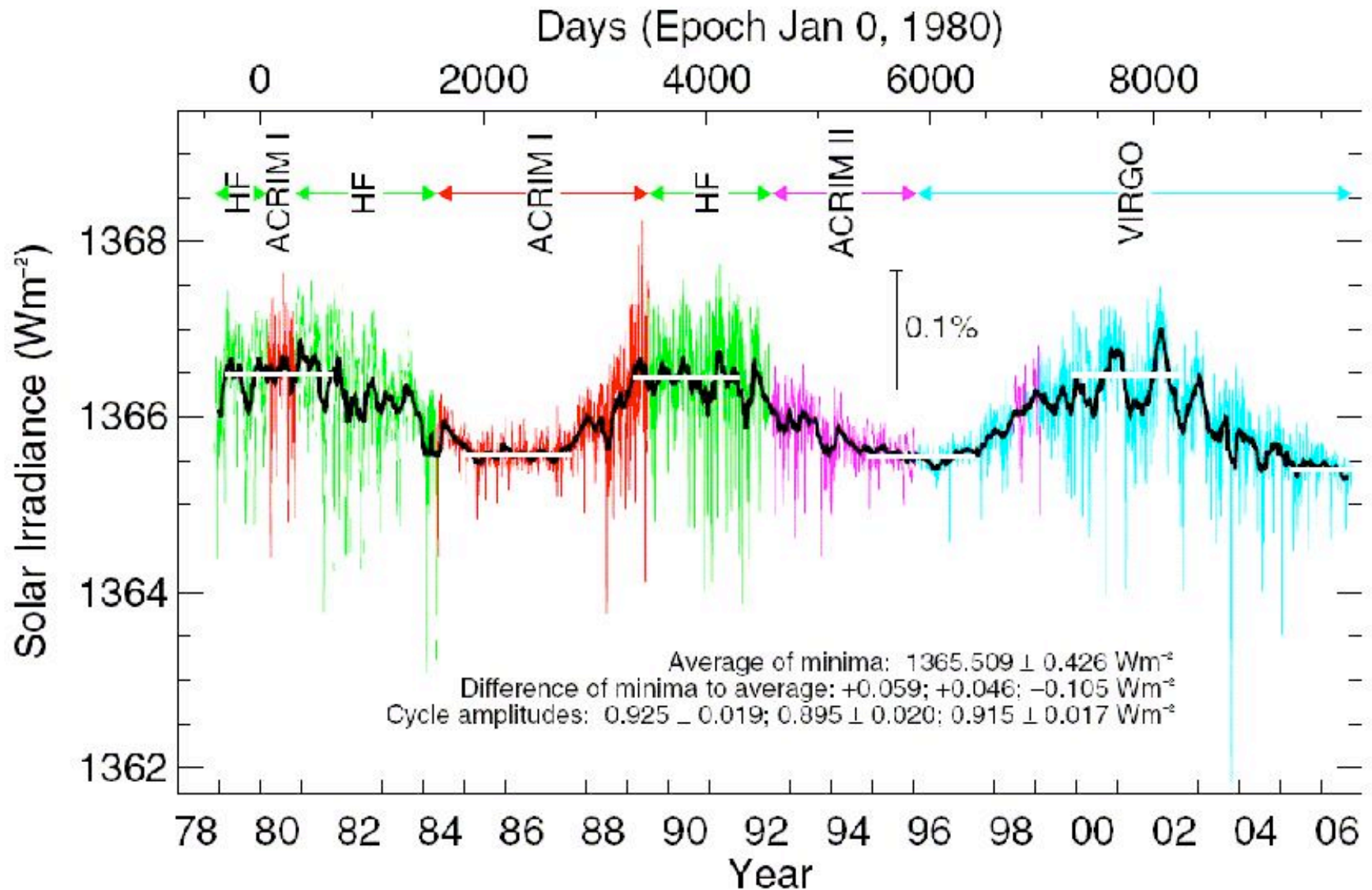
Energy radiated by the Earth

Energy received by the Earth

$$T_e = \sqrt[4]{\frac{S}{\sigma 4} (1 - A)}$$

(Equations found on page 43)

Satellite data (different satellites over time)



Planetary energy balance: E absorbed = E emitted

Combined equation: $\sigma T^4 = (S/4)(1-A)$ (p. 43)

$$S = 1370 \text{ W/m}^2$$

$$A = 0.3$$

$$\sigma = 5.67 \times 10^{-8} \text{ W/m}^2/\text{K}^4$$

**Substituting these values into equation, get
 $T_e \sim 255\text{K}$ (-18°C). (No greenhouse effect)**

$$\Delta T_g = T_s - T_e = 288 \text{ K} - 255 \text{ K} = 33^\circ\text{C}$$

If Earth had no clouds, then $A = 0.1$

If no clouds (and no greenhouse), what is T_e ? 271.4 K

Composition of the atmosphere; N₂, O₂, Ar are not greenhouse gases

Table 3-2 Major Constituents of Earth's Atmosphere Today

Name and Chemical Symbol	Concentration (% by volume)
Nitrogen, N ₂	78
Oxygen, O ₂	21
Argon, Ar	0.9
Water vapor, H ₂ O	0.00001 (South Pole)–4 (tropics)
Carbon dioxide, CO ₂	0.039 [*]

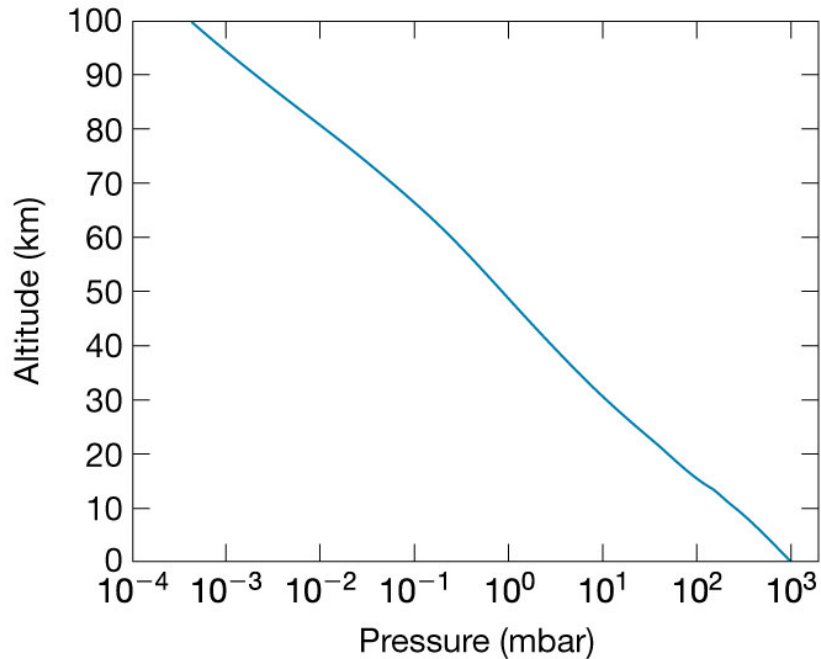
^{*}In 2008.

Table 3-3 Important Atmospheric Greenhouse Gases

Name and Chemical Symbol	Concentration (ppm by volume)
Water vapor, H ₂ O	0.1 (South Pole)–40,000 (tropics)
Carbon dioxide, CO ₂	390
Methane, CH ₄	1.7
Nitrous oxide, N ₂ O	0.3
Ozone, O ₃	0.01 (at the surface)
Freon-11, CCl ₃ F	0.00026
Freon-12, CCl ₂ F ₂	0.00048

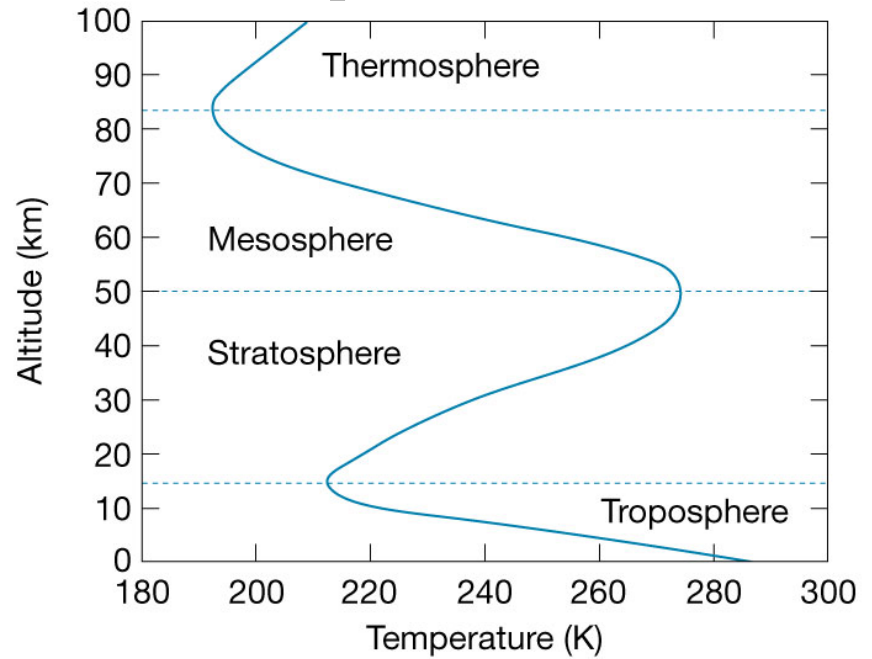
Structure of the atmosphere

P decreases exponentially
with increased altitude



(a)

T decreases then increases
from surface up through
stratosphere



(b)

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Fig. 3-9

*But why does T
increase in stratosphere?*

Why does T go up in stratosphere at ~50 km?: Ozone absorbs solar UV radiation. (Ozone conc. is highest at 30 km, but more UV is available at 50 km.) This means ozone is a greenhouse gas.

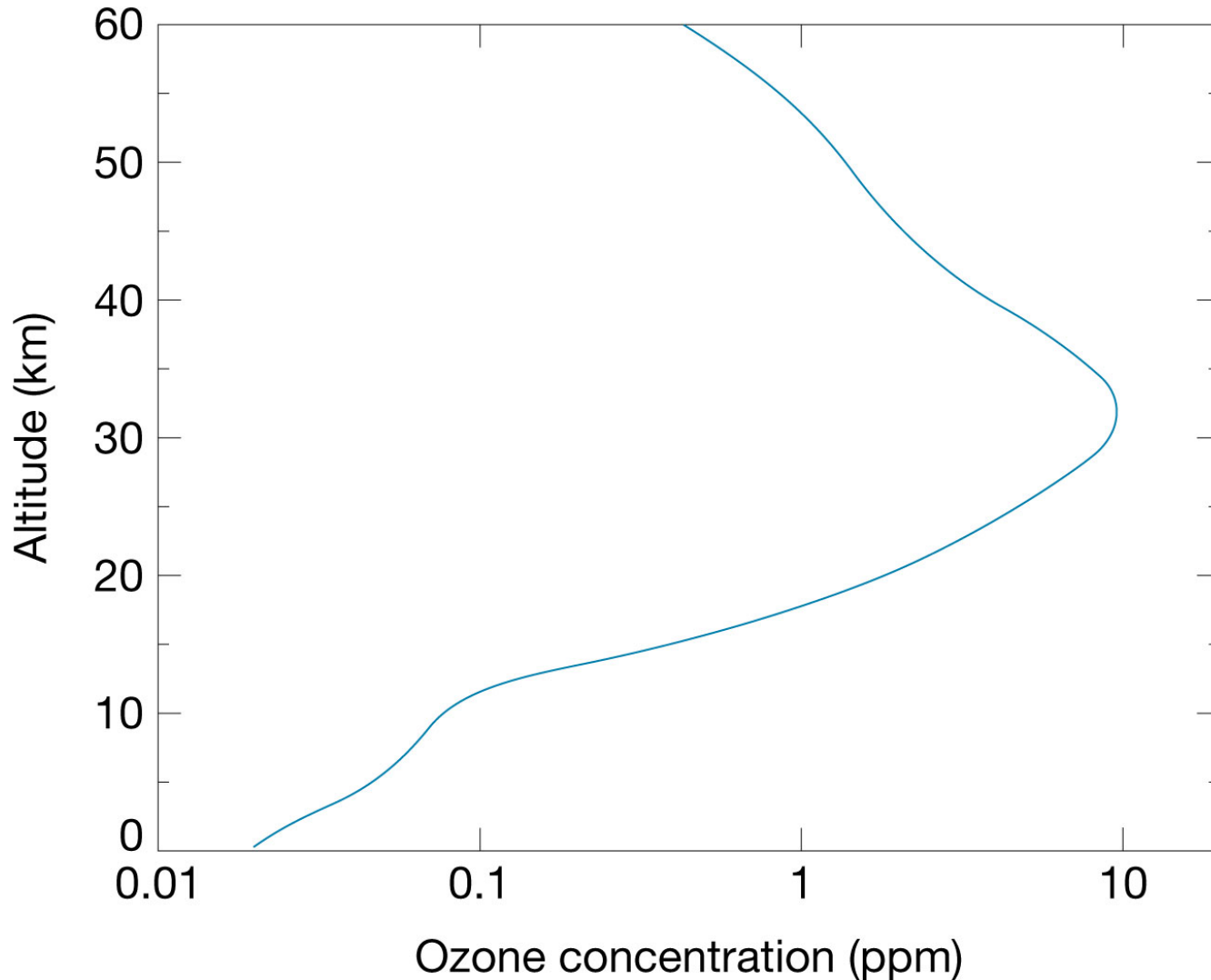
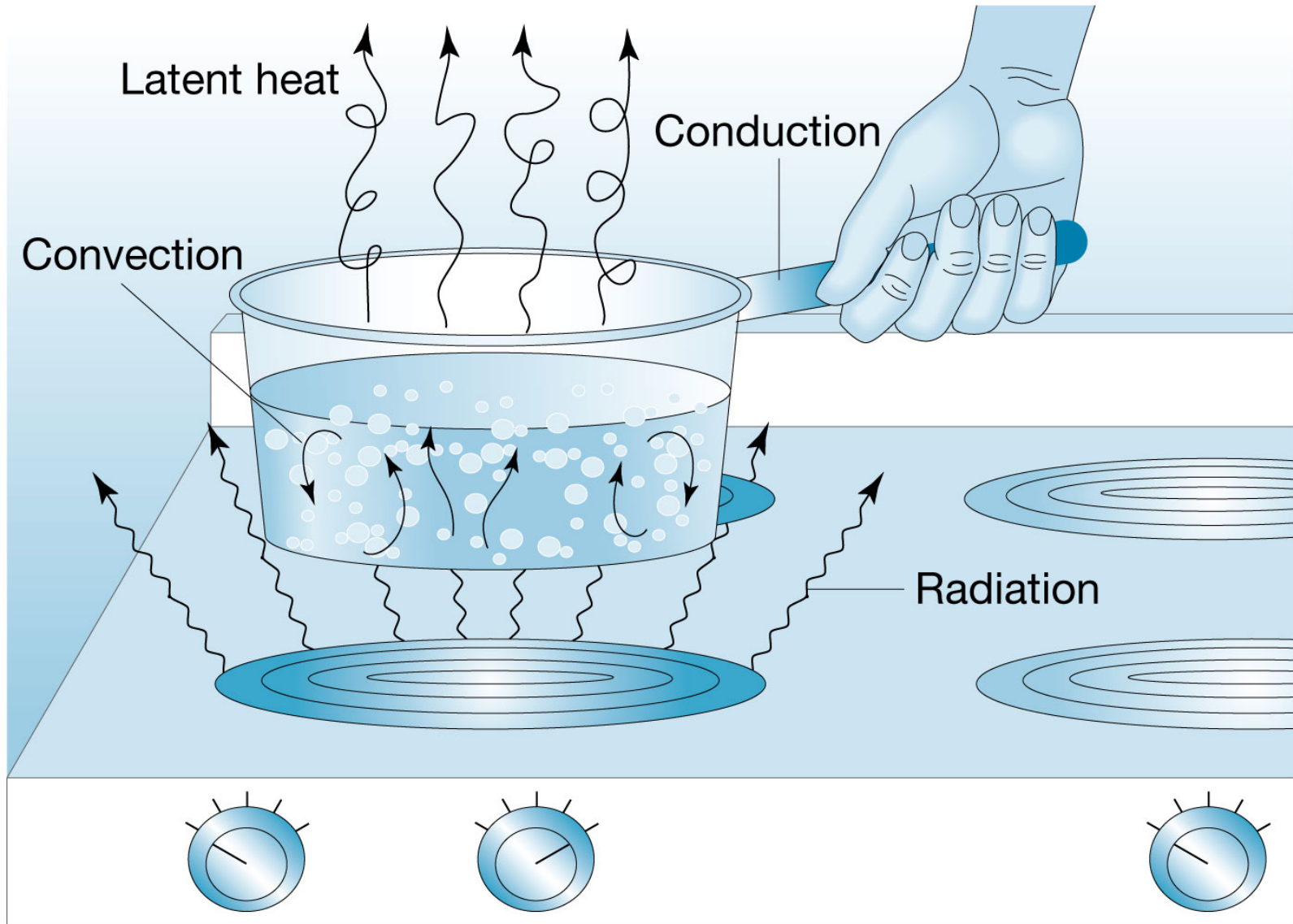


Fig. 3-11

Fig. 3-10

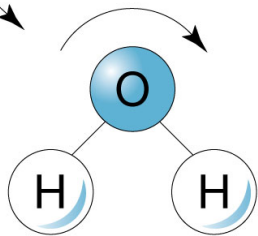
Heat Transfer...



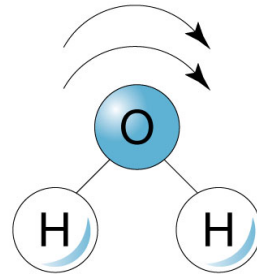
Greenhouse gas: must absorb & emit IR radiation

Incoming IR photon

Fig. 3-12



Slow rotation rate



Faster rotation rate

If e-m wave has just the Right frequency, molecule can absorb it, which makes the molecule rotation rate increase. The frequency needed depends on the molecule's structure.

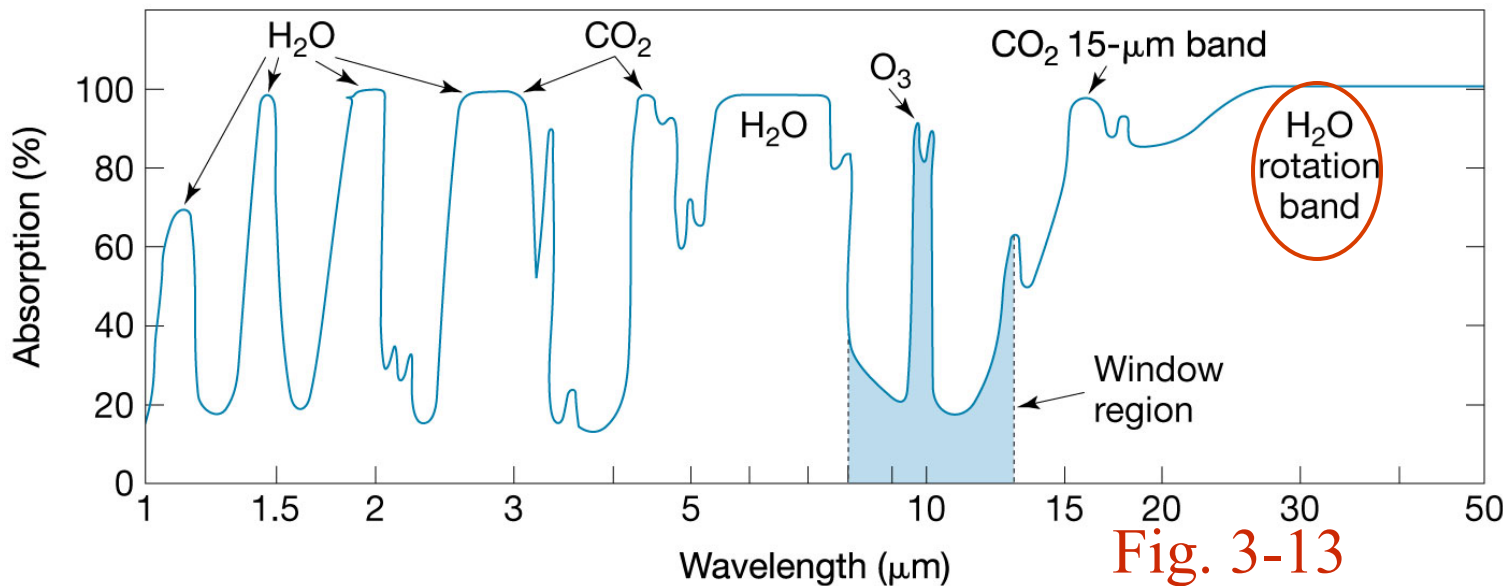
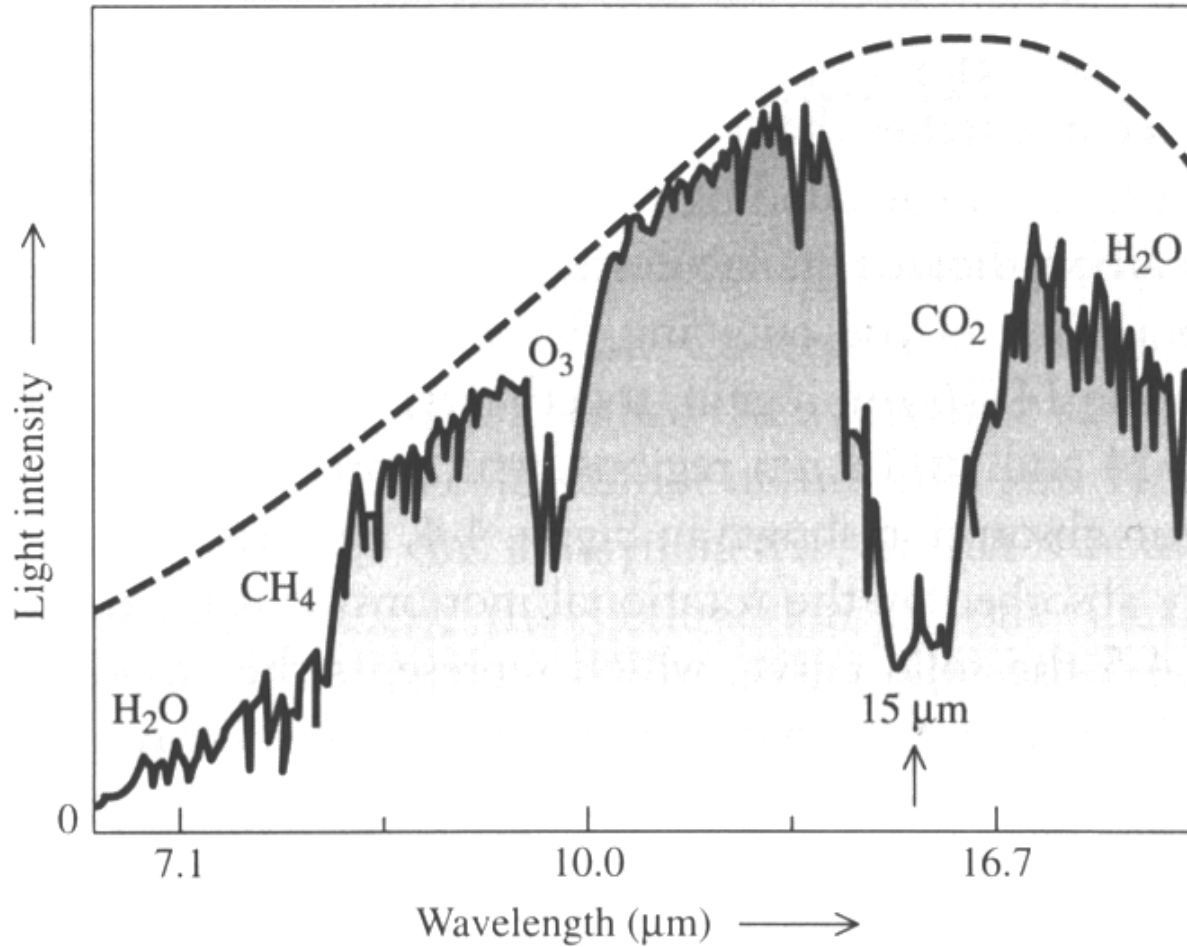
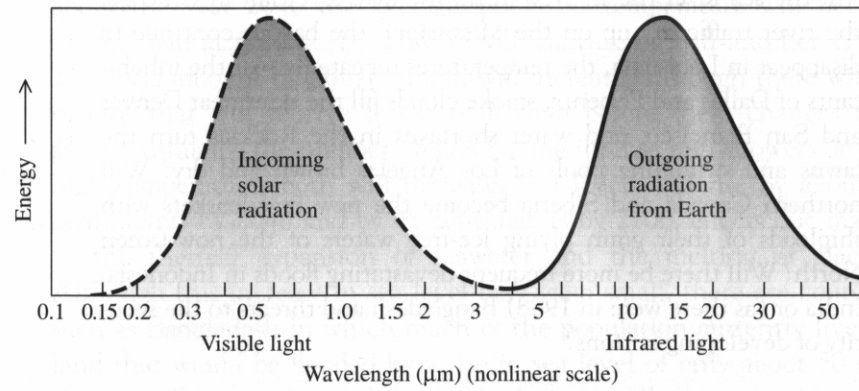
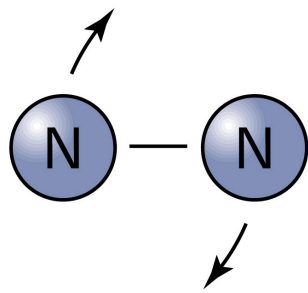


Fig. 3-13

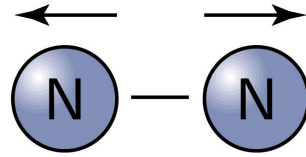
Water absorbs E with 12 to 1000 microns λ (microwaves)





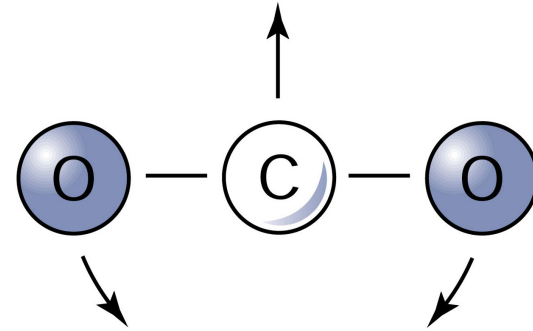
Rotation

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Vibration

Fig. 3-14,15

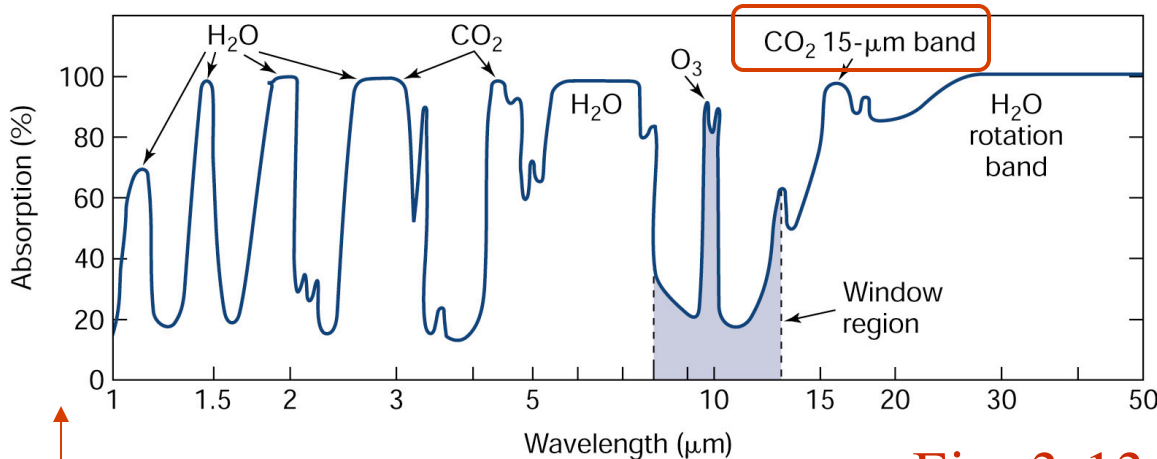


Bending mode
(15- μm band)

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If frequency of wave matches frequency of vibration, photon can be absorbed and molecule vibrates more vigorously.

A CO₂ molecule can vibrate in several ways including by bending. The bending mode has a 15 micron wavelength.

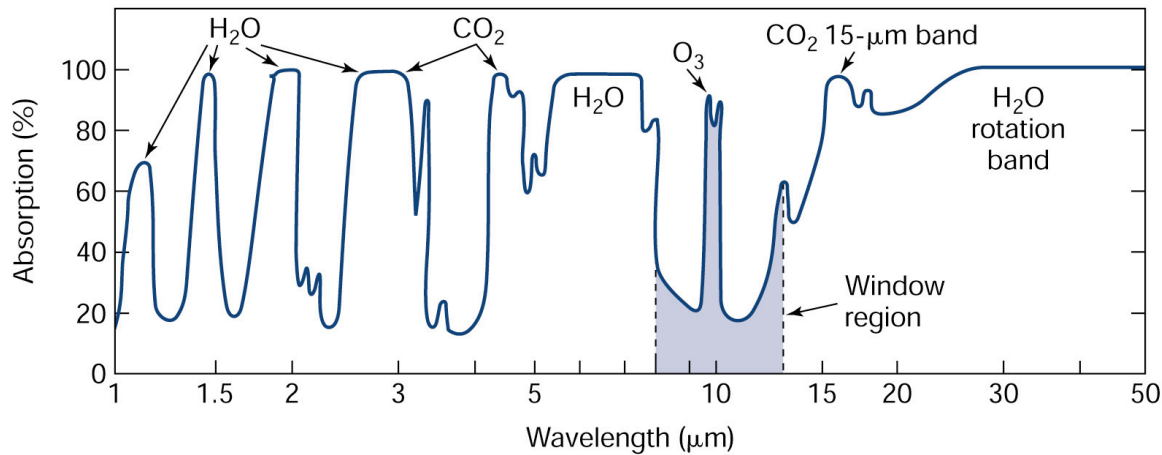


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Fig. 3-13

Visible light

Earth's IR emissions peak at ~ 15 microns, which is why CO₂ is such an important greenhouse gas.



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Fig. 3-13

“Window region” is where H₂O and CO₂ are poor absorbers. Freons absorb in this frequency range, so does ozone.

On a separate note...

N₂ and O₂ are perfectly covalent and have no separation of positive and negative charge in the molecule (no dipole); the molecule is also symmetrical. An e-m wave can't have a *net* interaction with a totally symmetric molecule.



(a)

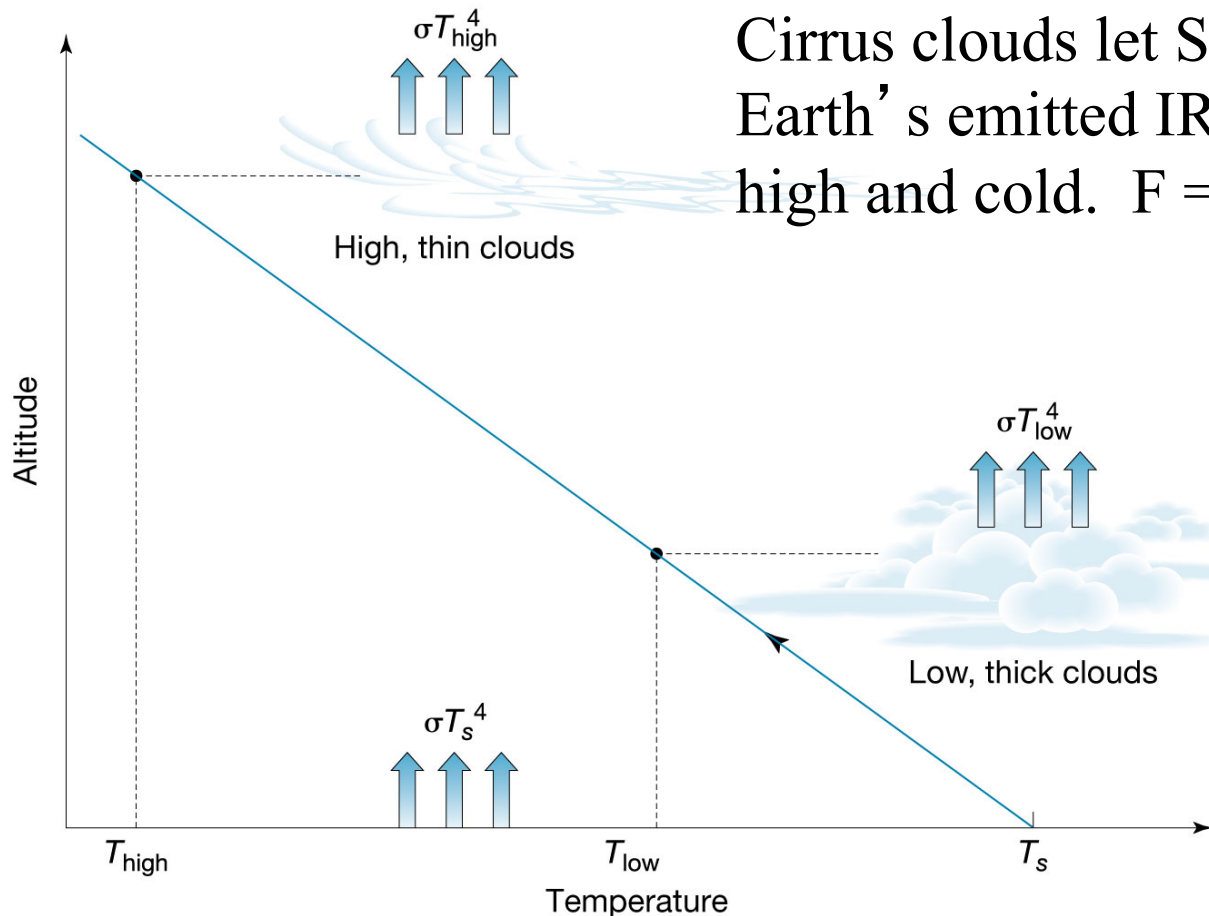
Fig. 3-17a



(b)

Fig. 3-17b

Clouds: they are white, and increase Earth's albedo
 They are made of H₂O and absorb some IR radiation emitted by Earth. They both cool and warm!



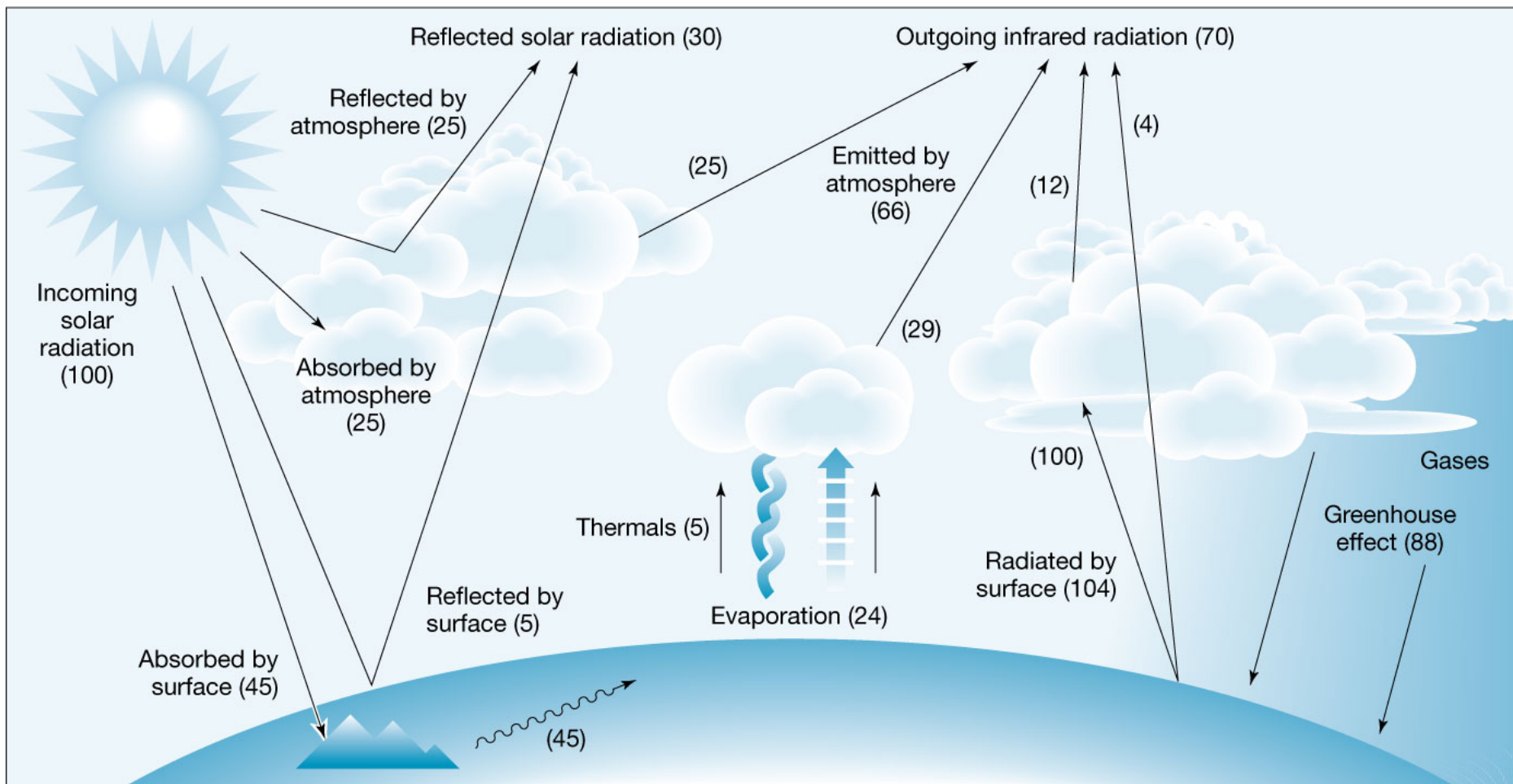
Cirrus clouds let Sun E through, but absorb Earth's emitted IR radiation. They are up high and cold. $F = \sigma T^4$

Stratus clouds reflect Sun's rays. They are warmer, so they radiate more IR energy to space.

Fig. 3-18

Clouds are tricky in climate models!

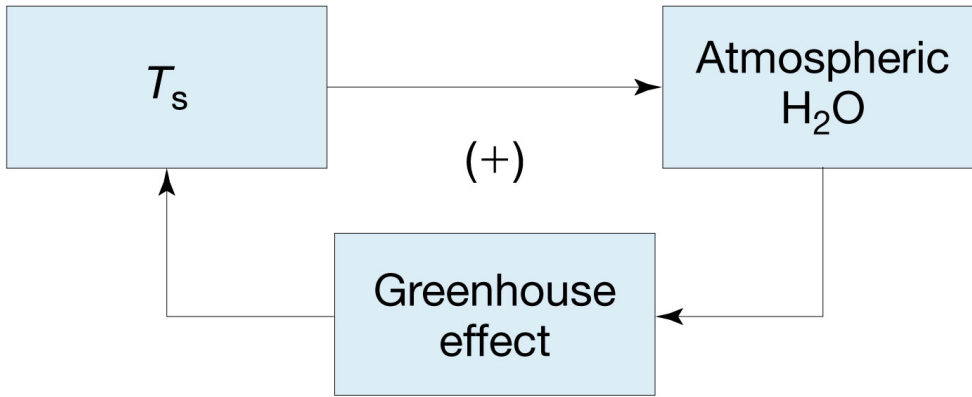
Earth's atmospheric energy budget, setting solar radiation to 100



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Fig. 3-19

Climate feedbacks:

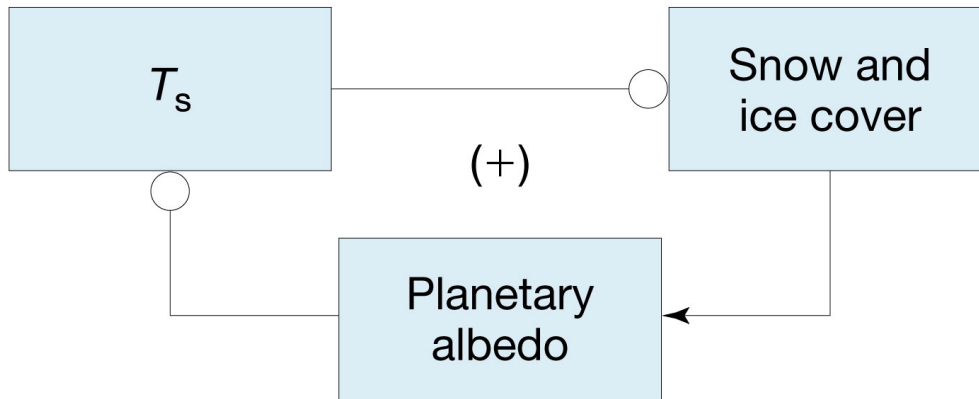


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Fig. 3-20

Water vapor: if T goes up
(maybe due to more CO₂)...

(this pos. fb. doubles effect
of surface T increase in
response to CO₂ increase)

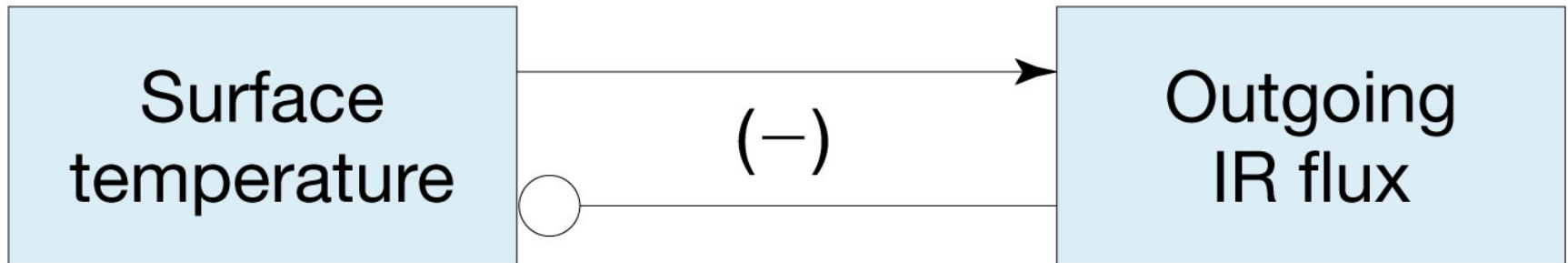


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Fig. 3-21

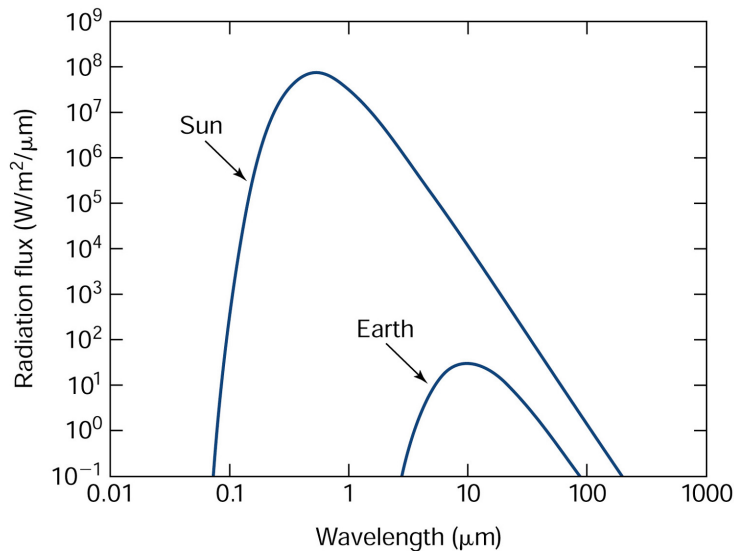
Snow and Ice: if T goes down...

We'd be in trouble if it weren't for IR flux/T feedback!



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Fig. 3-22



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Stefan-Boltzmann law:
 $F = \sigma T^4$, where σ is a constant

Fig. 3-8

Global E balance

- Earth is warmed by absorption of visible light from the Sun and cooled by emission of IR radiation to space
- Greenhouse gases absorb IR radiation by increasing rotation or vibration rates. Each gas absorbs particular wavelengths
- Feedbacks regulate climate change:
 - Water vapor, snow and ice positive fb loops
 - IR flux-T negative feedback loop
 - Clouds contribute to climate feedback but effect is not well known.